



# Assimilation of GPM/DPR at JMA

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Data Assimilation Seminar in RIKEN/AICS, Kobe, Japan, 13 Jun, 2016

# OUTLINE

1. Introduction
2. Operational NWP system at JMA
3. DPR data assimilation
  - Radar simulator
  - 1D+4DVAR
    - Relative humidity estimation
  - Quality control
4. Performance evaluation
5. Summary

# Introduction

- JMA has been operating NWP models for weather forecasting and disaster prevention information providing.
- To make the initial condition of NWP model, JMA assimilates many observation data. Especially, **satellite data** are most important data for improvement of the initial condition.
- Impact of GPM/DPR data assimilation at JMA
  - GPM/DPR was started to assimilate operationally in March 2016.

# GPM core satellite

## Global Satellite Mapping of Precipitation Global (GPM core)

- **Dual-frequency Precipitation Radar (DPR)**
  - Japan Aerospace Exploration Agency (JAXA)
  - National Institute of Information and Communications Technology (NICT)

② マイクロ波放射計 (GMI)

GPM microwave imager

③ 太陽電池パドル

① 二周波降水レーダ (DPR)

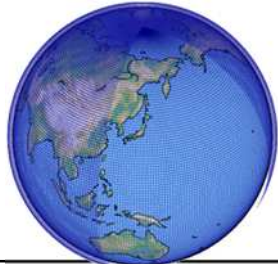
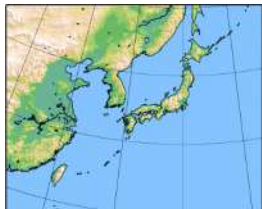


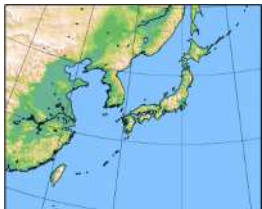
The DPR consists of a Ku-band precipitation radar (KuPR) and a Ka-band precipitation radar (KaPR).

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# Current NWP models of JMA

	In Operation					In Test Operation (*)
	Deterministic					
	Global Spectral Model <b>GSM</b>	Meso-Scale Model <b>MSM</b>	Local Forecast Model <b>LFM</b>	One-week Ensemble <b>WEPS</b>	Typhoon Ensemble <b>TEPS</b>	Meso-scale Ensemble <b>MEPS</b>
objectives	Short- and Medium-range forecast	Disaster reduction Aviation forecast	Aviation forecast Disaster reduction	One-week forecast	Typhoon forecast	Uncertainty and probabilistic information of MSM
Forecast domain	Global 	Japan and its surroundings (4080km x 3300km) 	Japan and its surroundings (3160km x 2600km) 	Global 	Japan and its surroundings (4080km x 3300km) 	
Horizontal resolution	TL959(0.1875 deg)	5km	2km	TL479(0.375 deg)	5km	
Vertical levels / Top	100 0.01 hPa	48+2 21.8km	58 20.2km	60 0.1 hPa	48+2 21.8km	
Forecast Hours (Initial time)	84 hours (00, 06, 18 UTC) 264 hours (12 UTC)	39 hours (00, 03, 06, 09, 12, 15, 18, 21 UTC)	9 hours (00-23 UTC hourly)	264 h (00, 12 UTC) 27 members	132 h (00, 06, 12, 18 UTC) 25 members	39h 11 members
Initial Condition	Global Analysis (4D-Var)	Meso-scale Analysis (4D-Var)	Local Analysis (3D-Var)	Global Analysis with ensemble perturbations (SV)		Meso-scale Analysis with ensemble perturbations (SV)

# Main Operational Forecast model

## Global NWP System

Global Spectral Model (GSM)

Horizontal resolution: TL959 (0.1875 deg)

Global Analysis (GA): 4D-Var

## Meso-scale NWP System

Meso-scale model (MSM)

Forecast Model: JMA-NHM

Horizontal resolution: 5 km

Meso Analysis (MA): 4D-Var

Data assimilation system: JNoVA

## Local NWP System

Local Forecast model (LFM)

Forecast Model: ASUCA

Horizontal resolution: 2 km

Local Analysis (LA): 3D-Var

Analysis cycle

Data assimilation system: ASUCA-Var

# Meso-scale NWP System

- Main purpose: Providing disaster prevention information
  - Accuracy of precipitation forecast is of very importance.
  - Hydrometeors in initial condition must be improved for forecast improvement.
- Cloud microphysics process
  - 3-ice 6-class bulk scheme
  - Prognostic hydrometeors
    - Water vapor, cloud, rain, ice, snow and graupel
  - Reflectivity calculation needs these hydrometeors in data assimilation system.

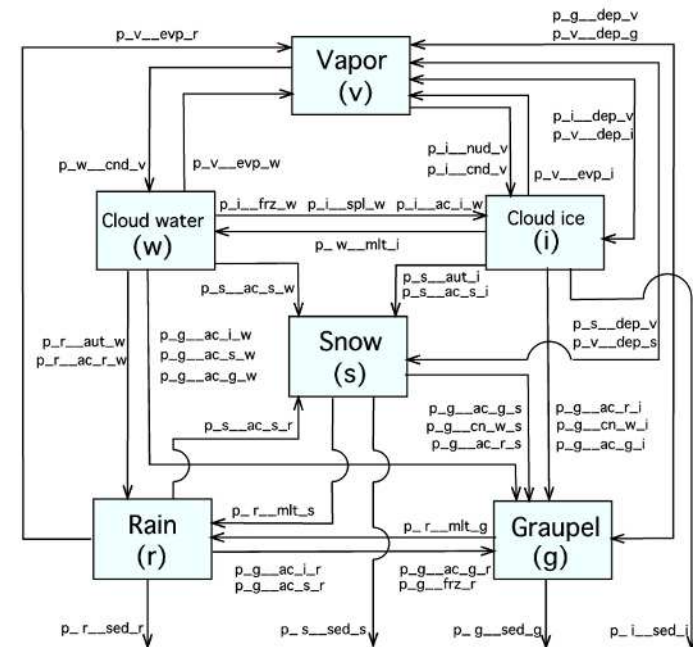
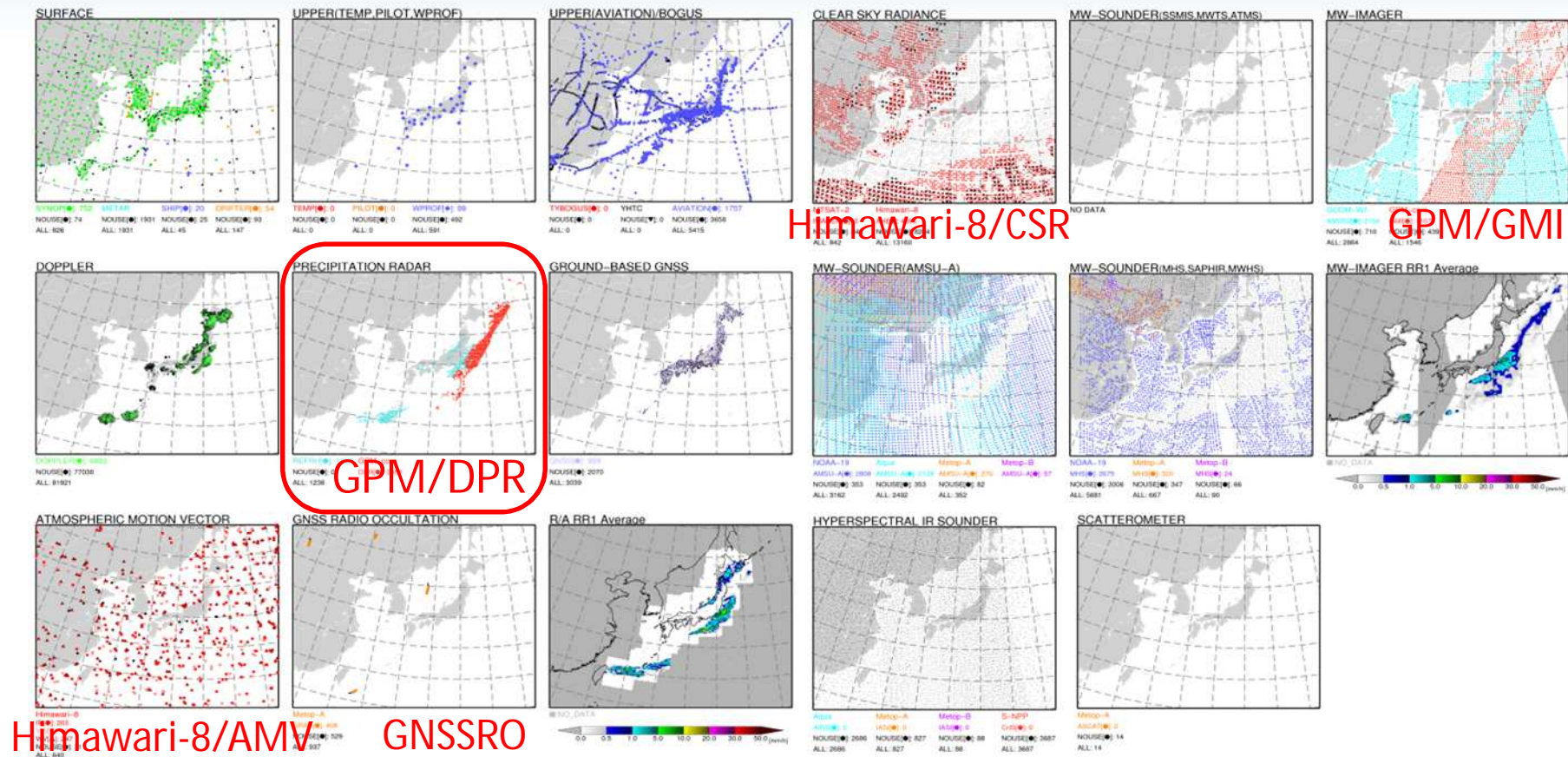


Figure 3.5.2: Cloud microphysical processes in the MSM. For a list of symbols, see Table 3.5.1.



# Coverage map of assimilated observation in Meso-scale NWP system

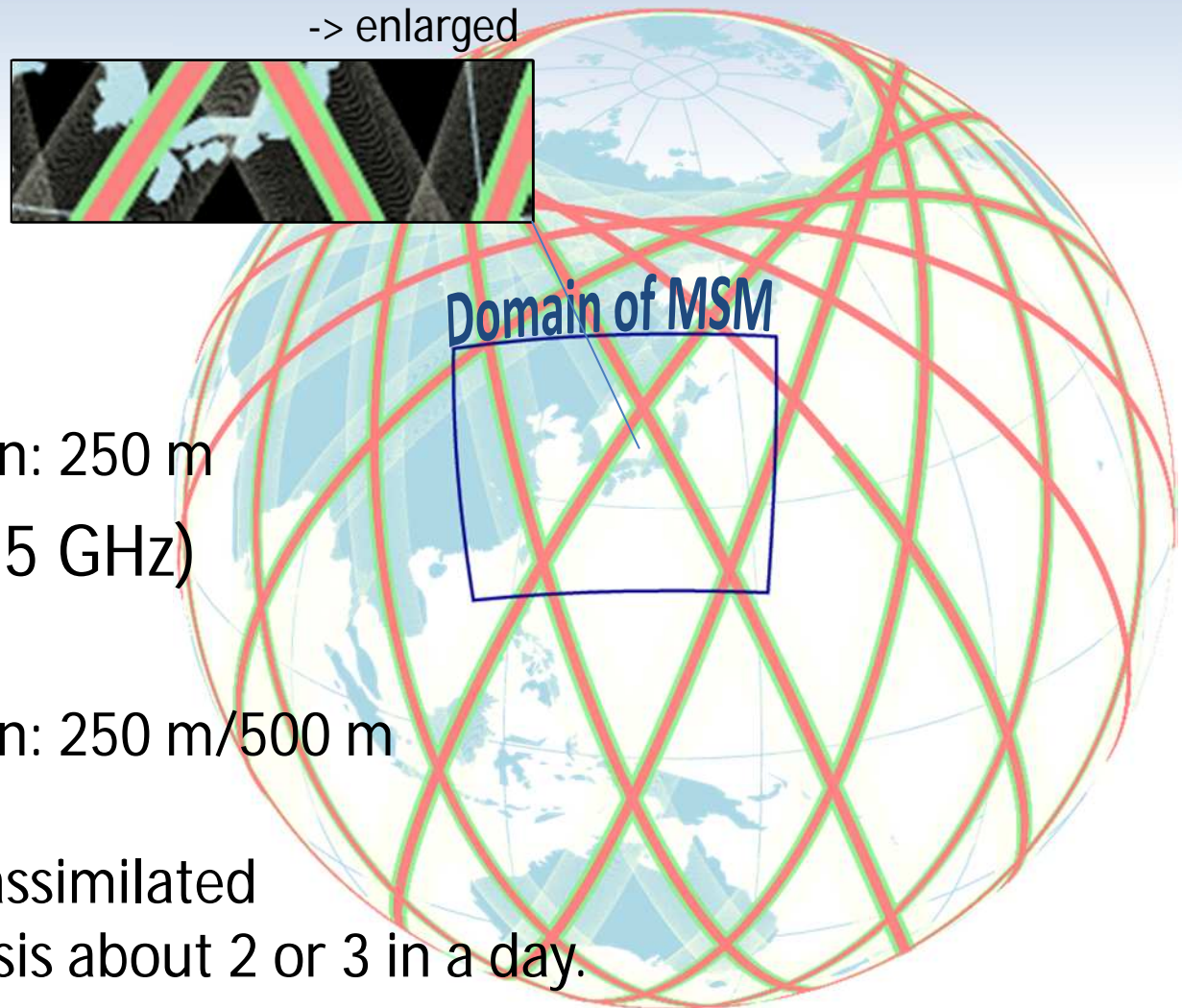


Recent updates: GPM/DPR, GPM/GMI, GNSSRO, Himawari-8/AMV, Himawari-8/CSR are started to assimilate operationally in March 2016.

# GPM Data Coverage during 24-hour

- GMI
  - Width: 800 km
- KuPR(13.6 GHz)
  - Width: 245 km
  - Vertical resolution: 250 m
- KaPR MS/HS\*(35.5 GHz)
  - Width: 125 km
  - Vertical resolution: 250 m/500 m

\* High sensibility mode
- KuPR and KaPR are assimilated by meso-scale analysis about 2 or 3 in a day.



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# Ground-based Radar Simulator in JMA

- It was developed for weather radar assimilation.

## ▣ Beam bending

- Refraction index

$$N = \frac{77.6}{T} \left( p + 4810 \frac{e}{T} \right)$$

- The Earth curvature

## ▣ Effective hydrometeors

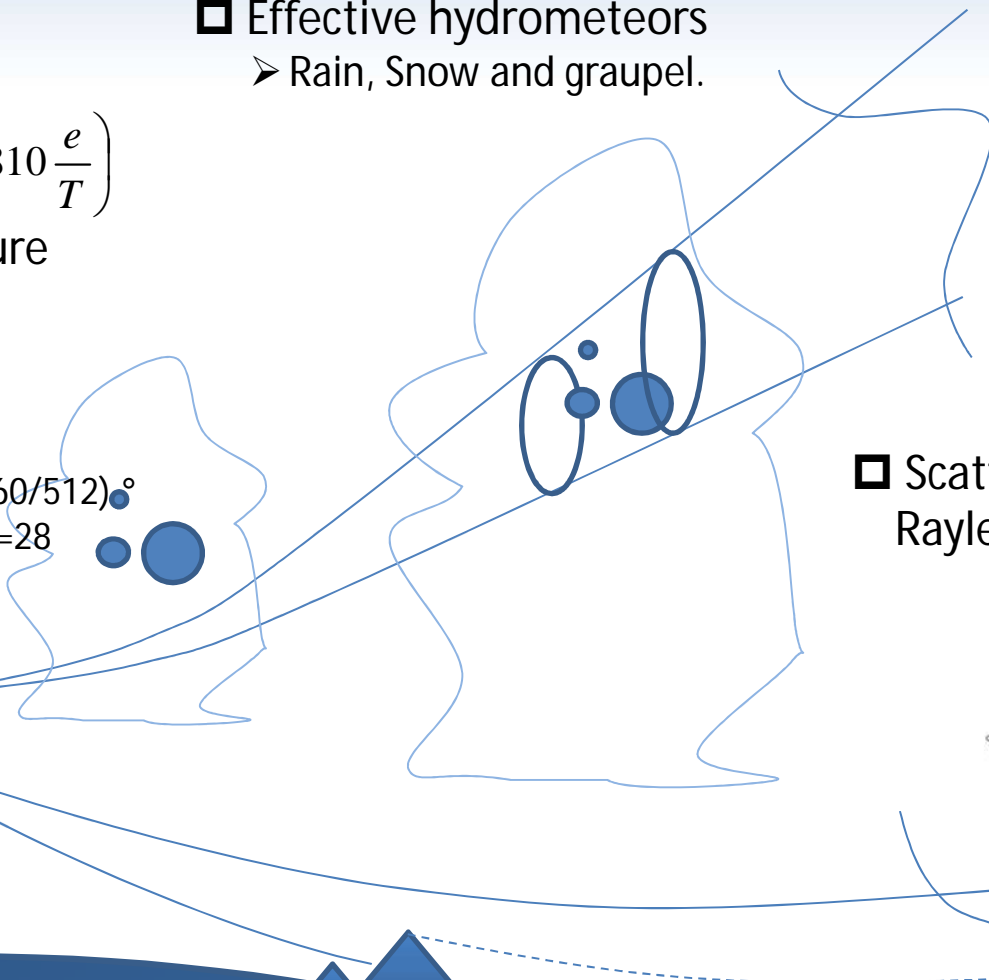
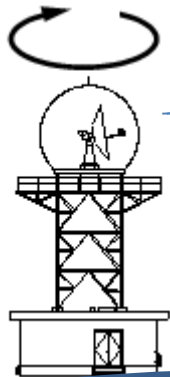
- Rain, Snow and graupel.

## ▣ Beam shape

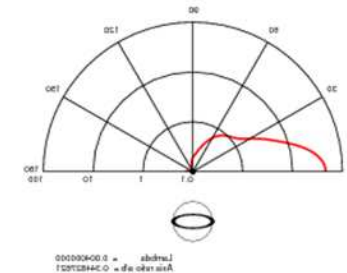
- Gauss-Hermite quadrature, quadrature order n = 5

## ▣ Virtual radar site

- Ex. Tokyo
- $\lambda = 5.7\text{cm}$
- Resolution =  $500\text{m} \times (360/512)^\circ$
- Number of elevation = 28



## ▣ Scattering cross-section Rayleigh (or T-matrix)



## ▣ Beam blockage by topography

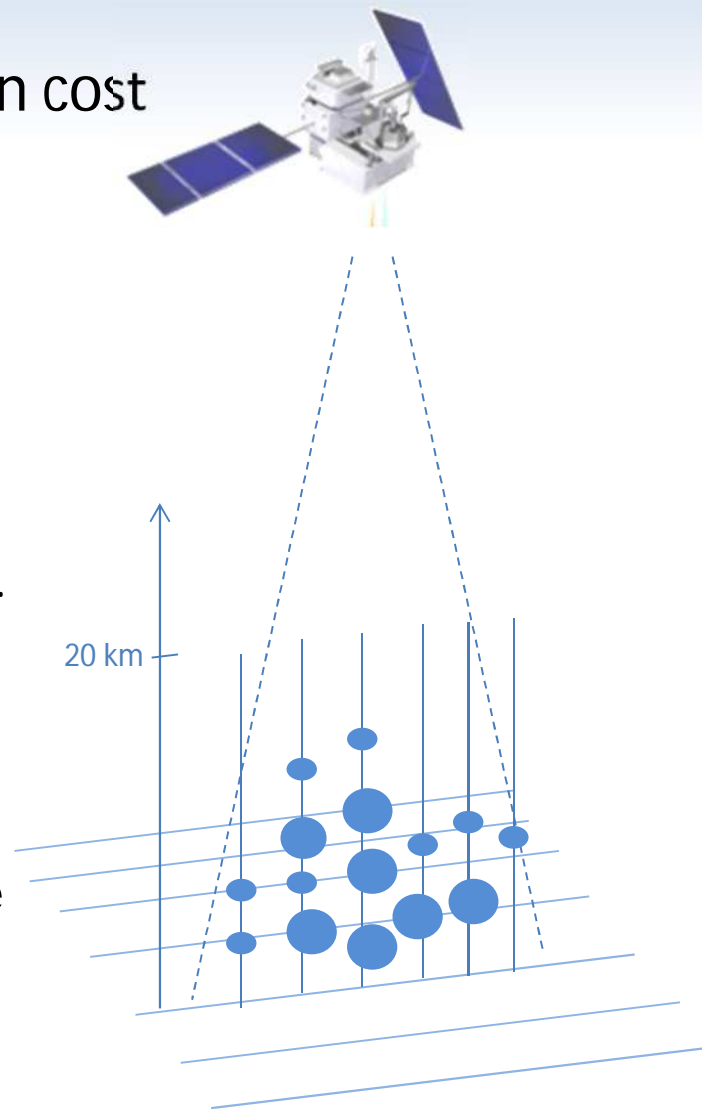
# Space-borne radar simulator

- Simplification for reducing computation cost

- Slant beam path, beam width and beam bending
  - Horizontal resolution of MSM is 5km, it is coarse-grid about beam simulation.
  - Small Impact for DA
- Attenuation
  - Corrected Z factor products has been used.

- For computing efficiency

- Z factor Table is prepared at offline.
- In online, Z factor is given by Look up table method



# Space-borne radar simulator

- Reflectivity calculation

- Effective particle: Rain, snow and graupel
  - Cloud water and cloud ice are ignored.
- Size distribution: Negative exponential dist.
  - Intercept parameter is fixed
- Particle shape: sphere
- Scattering calculation: Lorenz-Mie theory
  - Single scattering

$$\eta_x = \int_{D_{\min}}^{D_{\max}} \sigma_{bx}(D) N_x(D) dD$$

- Dielectric constant

- Water: Debye
- Snow: Boren and Batton (1982)

- Size distribution:

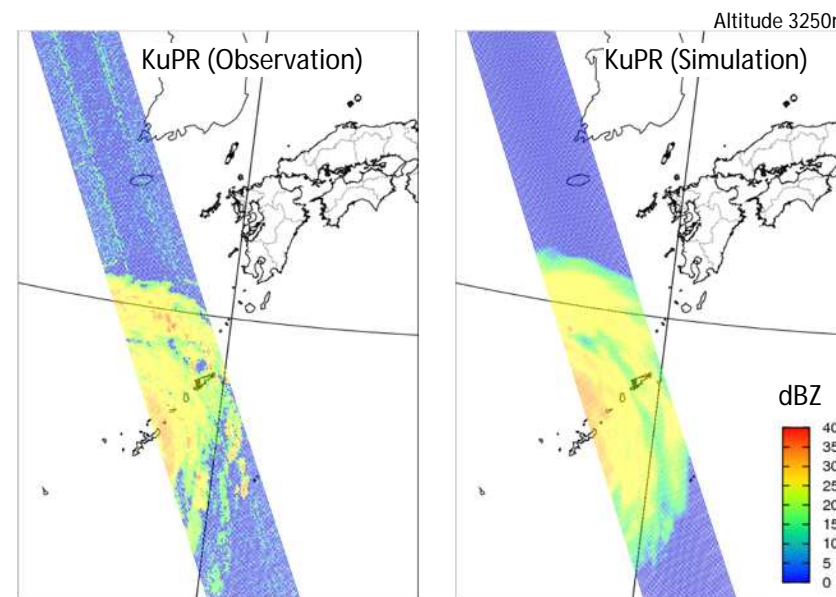
$$N_x(D) = N_{x0} \exp(-\lambda_x D_x)$$

- Slope parameter:

$$\lambda_x = \left( \pi \frac{\rho_x N_x}{\rho_a Q_x} \right)^{\frac{1}{3}}$$

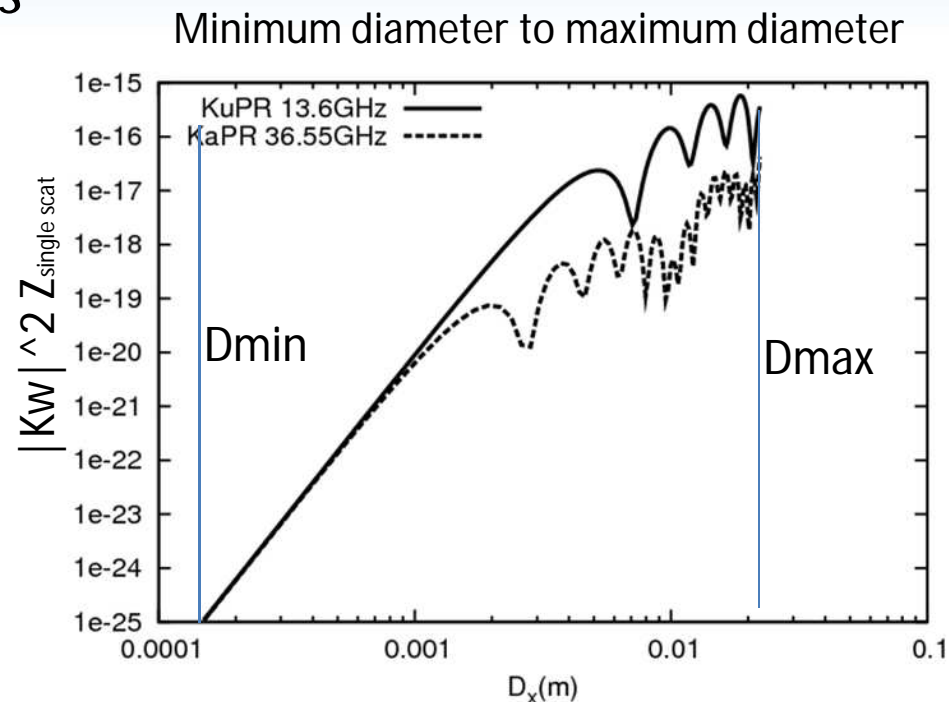
- Intercept parameter:

$$N_{x0} = N_x \lambda_x$$



# Space-borne radar simulator

- Bias of simulated Z factor exists in ice phase (Eito and Aonashi 2009)
  - Case of fixed intercept parameter in 1-moment scheme
    - Insufficient to describe size distribution
    - Especially, error becomes large at large diameter.
  - Case of unfixed intercept parameter in 2-moment scheme
    - Better scheme to describe size distribution than 1-moment scheme
- Operational model
  - 1-moment scheme
    - > Large bias caused by error of large size particle

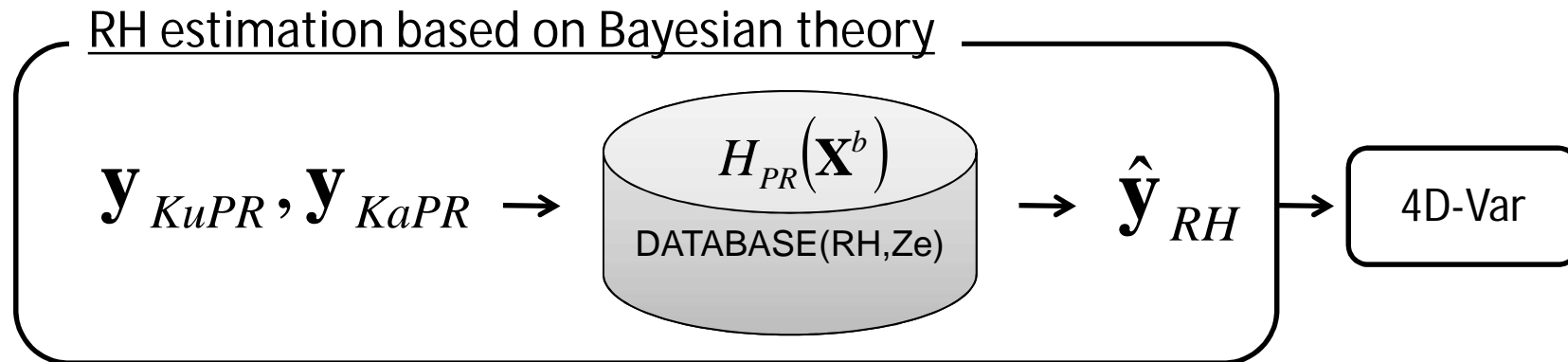


## Comparing the KuPR and KaPR

KuPR > KaPR from large size particles.  
Model bias affects simulated KuPR than KaPR.

# GPM/DPR data assimilation

- Assimilation method of KuPR and KaPR
  - 1D+4D-Var method
    - This method is same as ground based radar assimilation at JMA.(Ikuta and Honda, 2011)
  - 1. RH is retrieved from observed reflectivity, simulated reflectivity and first-guess. (Caumont et al., 2010)
  - 2. This retrieved RH is assimilated in the same way as conventional data by 4D-Var.





# Bayesian Theory

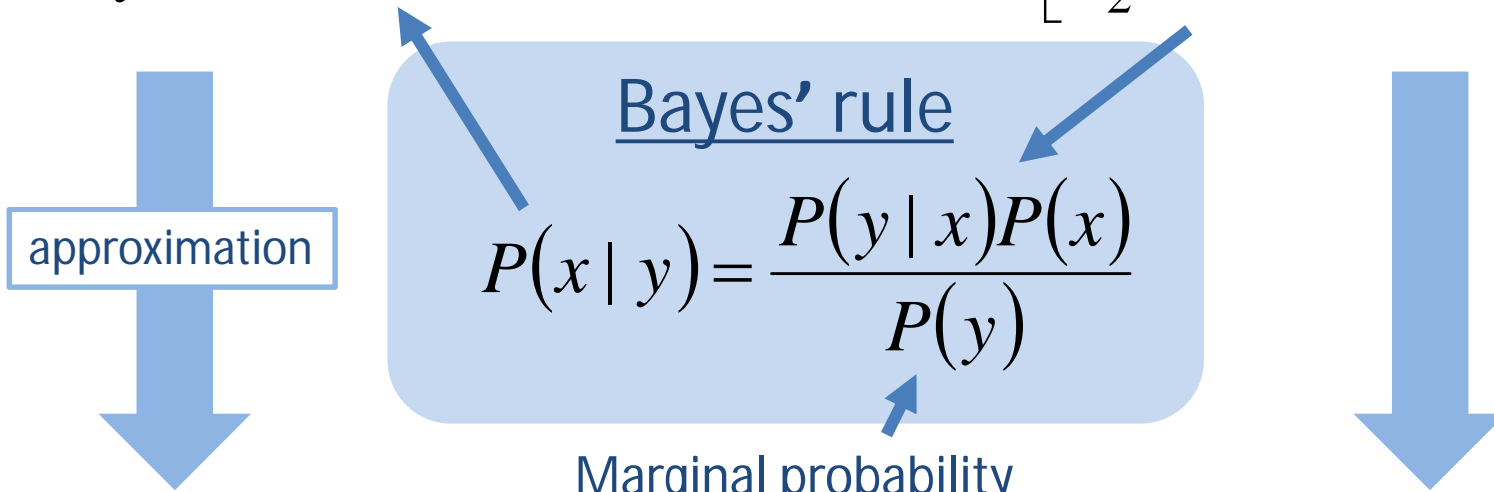
- RH estimation is based on Bayesian theory.

## Best estimation

$$\hat{x} = \int x \cdot P(x | y) dx$$

## Conditional probability

$$P(y | x) \propto \exp\left[-\frac{1}{2}(H(x) - y)^T R^{-1}(H(x) - y)\right]$$



approximation

## Bayes' rule

$$P(x | y) = \frac{P(y | x)P(x)}{P(y)}$$

## Marginal probability

$$P(y) = \int P(y | x)P(x)dx$$

## Weighted average

$$\hat{x} \equiv \sum_i x_i \frac{w_i}{\sum_j w_j}$$

## Weight function

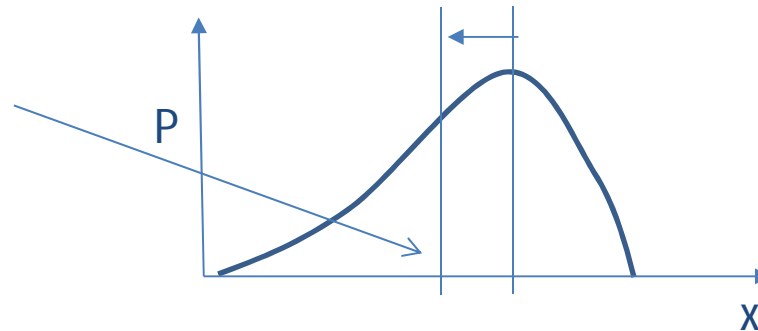
$$w_i = \exp\left[-\frac{1}{2}(H(x_i) - y)^T R^{-1}(H(x_i) - y)\right]$$

# Bayesian theory with Kernel density estimation

Weighted average  $\neq$  Maximum likelihood

→ Under estimation in case of non-Gaussian

$$\hat{x} = \sum_i x_i \frac{w_i}{\sum_j w_j}$$



*Our approach*

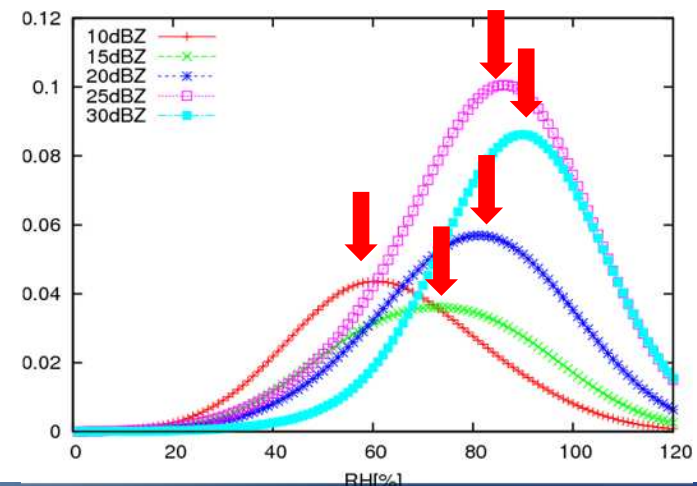
## Maximum likelihood using kernel density estimation

$$\hat{x} = \arg \max_x L(x | y_{Ku}, y_{Ka})$$

This method describes the likelihood function in a superposition of the Gaussian kernel.

$$L(x | y_1, y_2) \propto \frac{1}{nh_1h_2} \sum_i^n \left[ \prod_{j=1}^2 K\left(\frac{H(x_i) - y_j}{h_j}\right) \right]$$

Kernel function:  $K(t_i) = w_i \frac{1}{\sqrt{2\pi}} \exp(-t_i^2 / 2)$



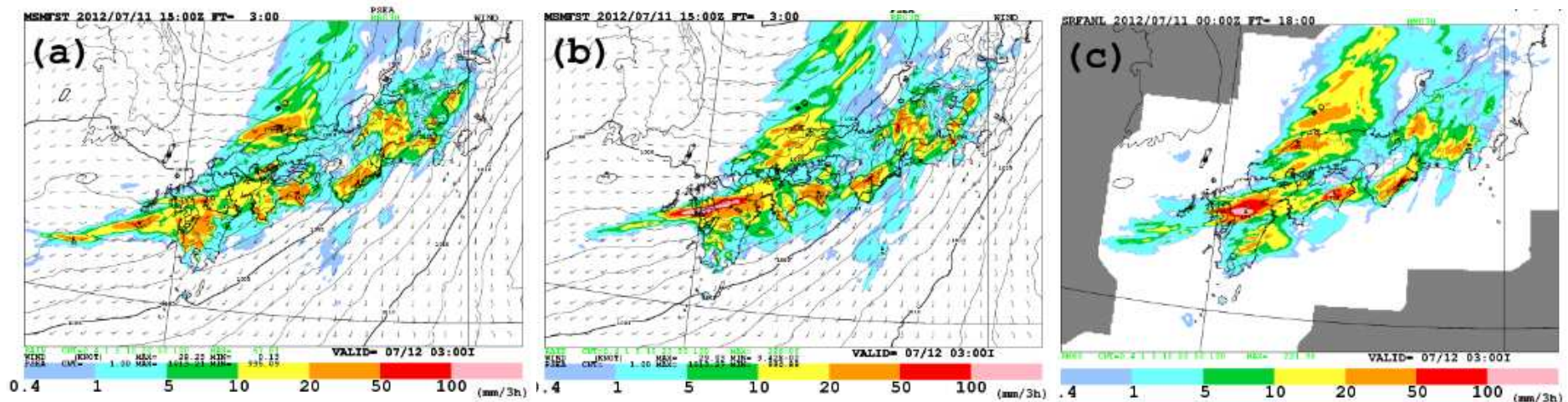
# Impact of weighted average v.s. kernel density

Assimilation experiment using ground-based weather radar of JMA

EXP: Weighted average

EXP: Kernel density

Observation



Precipitation forecast improved using the Kernel density approach than the weighted average approach.

# Estimation of relative humidity profiles

The RH profiles are updated using reflectivity observation profiles.

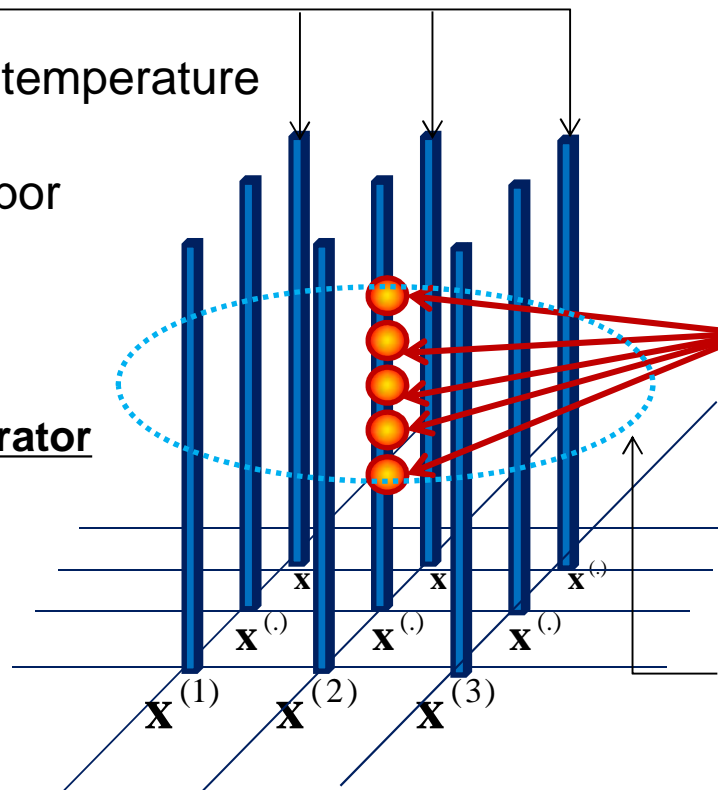
- The set of column in the first-guess are regarded as the pseudo ensemble members.
- Making a database of the relation between RH and Z.
- The RH profiles are estimated by based on Bayes' theorem with kernel density estimation.

## State vector

$$\mathbf{x} = \begin{pmatrix} \theta \\ p \\ q_v \\ q_r \\ q_s \\ q_g \end{pmatrix} \begin{matrix} \text{Potential temperature} \\ \text{Pressure} \\ \text{Water vapor} \\ \text{Rain} \\ \text{Snow} \\ \text{Graupel} \end{matrix}$$

## Observation operator

$$H_{DPR}(\mathbf{x})$$



## Observation vector

$$\mathbf{y} = \mathbf{Z} \leftarrow \text{KuPR, KaPR profile}$$

## Pseudo ensemble vector

$$\mathbf{X}^b = \left( \mathbf{x}^{(1)} \dots \mathbf{x}^{(m)} \right)$$

## Schur production of the observation

$$\text{localization } \mathbf{L}_{d,T}$$

# Bayesian theory with Kernel density estimation

## Adaptive bias correction

Bias is recalculated every analysis.

$$\langle d \rangle = \langle H(x) - y \rangle$$

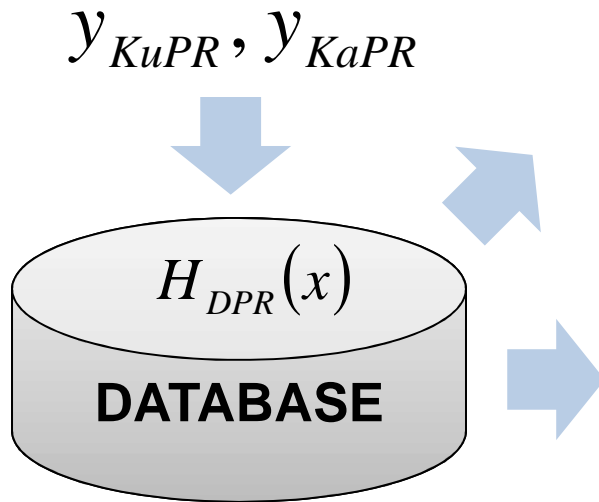
## Model probability density

$P(x|y)$  is redefined using  $\langle d \rangle$ .

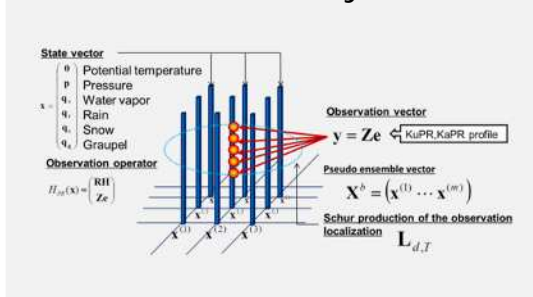
$$P(y|x) \propto \exp\left[-\frac{1}{2}(H(x) - y - \langle d \rangle)^T R^{-1}(H(x) - y - \langle d \rangle)\right]$$

## Maximum likelihood using Kernel density estimation

$$RH = \arg \max L(x | y_{Ku}, y_{Ka})$$

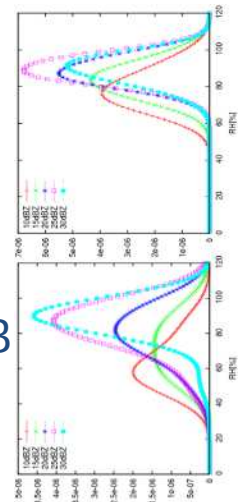


Database is made by model columns.

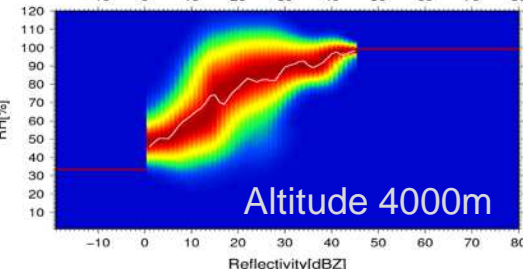
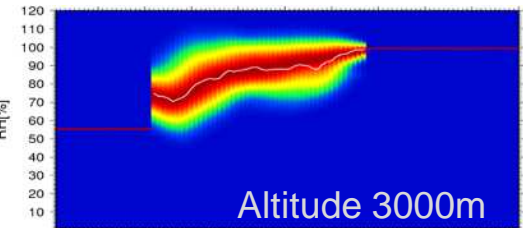


## 4DVar

Correlation of between B and R is neglected.

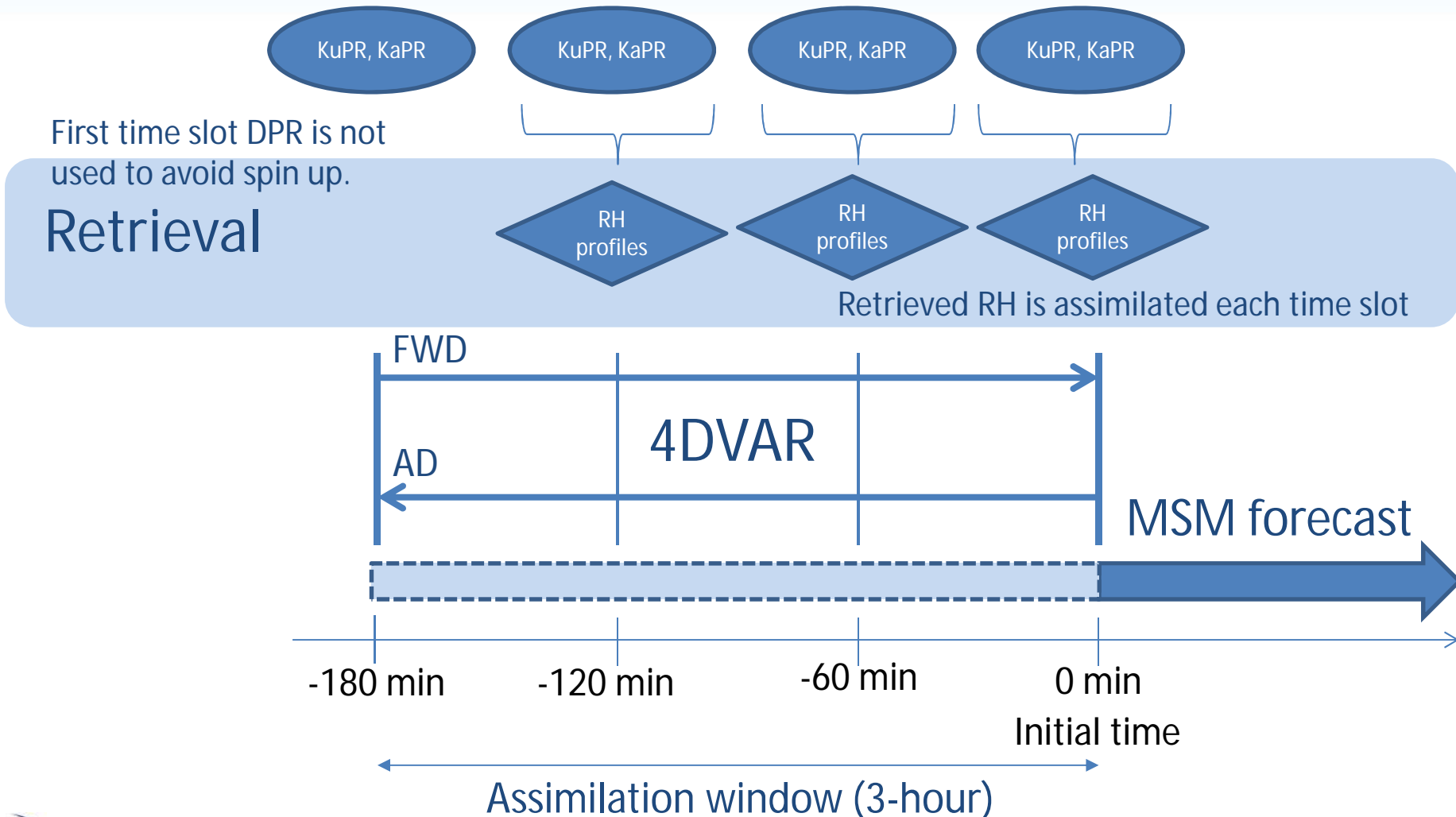


## Likelihood Function

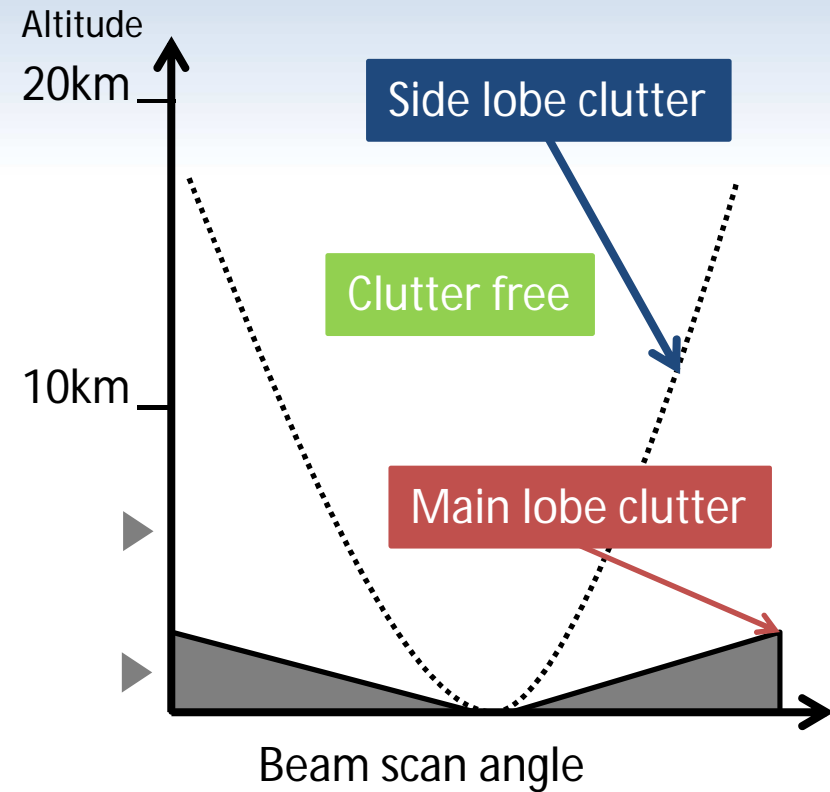
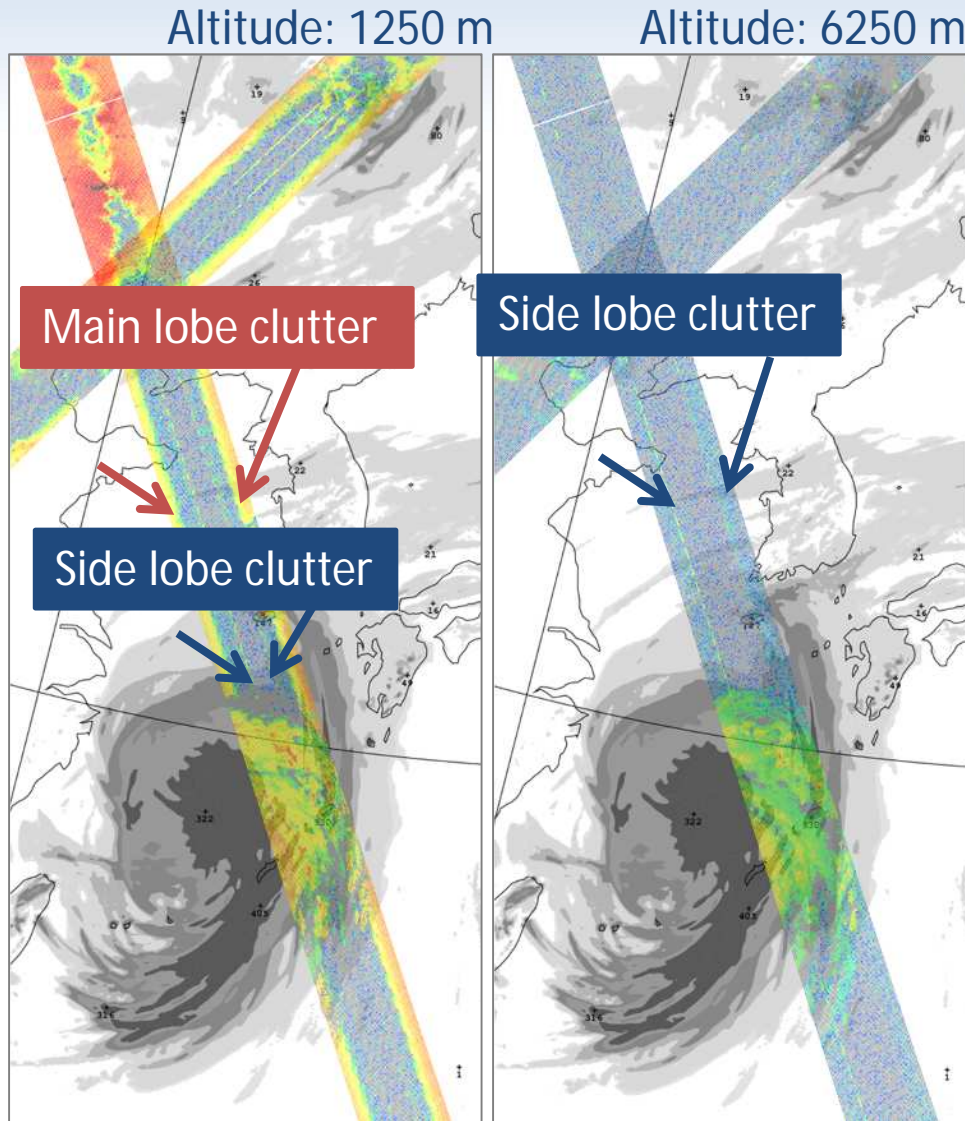


# DPR assimilation in Meso Analysis

- The observation are distributed to 4 time-slots by rounding off the observation time to hours.
- The observation within the period from 3.5 h before to 0.5 h after the initial time are assimilated.



# Quality Control [Clutter]

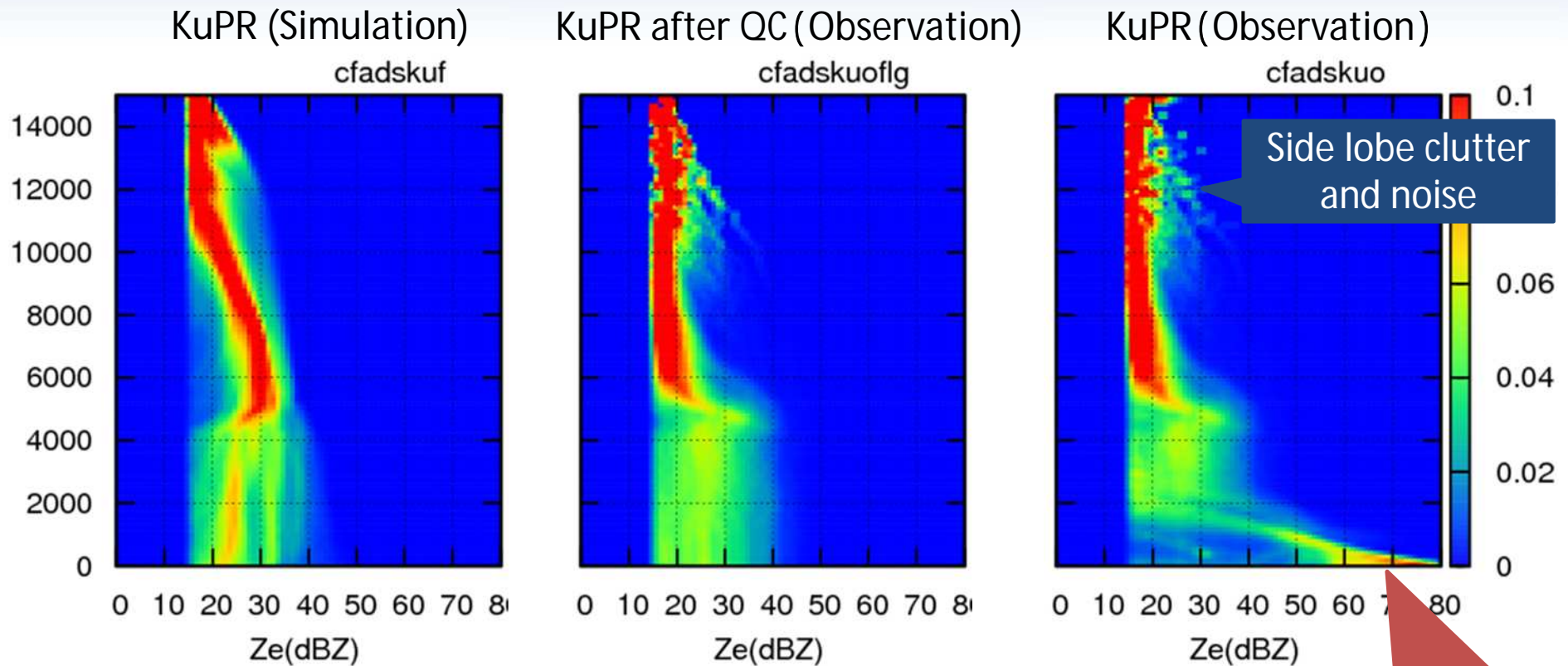


Ref. Tagawa et al. (2007)

- Value of both edge on the path is main lobe clutter.
- Two lines on the path are side lobe clutter.
- For DA, clutter must be removed.

# Quality Control [Clutter]

Comparisons between model and DPR using Contoured Frequency with Altitude Diagrams (CFADs)  
CFADs: > 15dBZ



Side lobe clutter and noise

Main lobe clutter

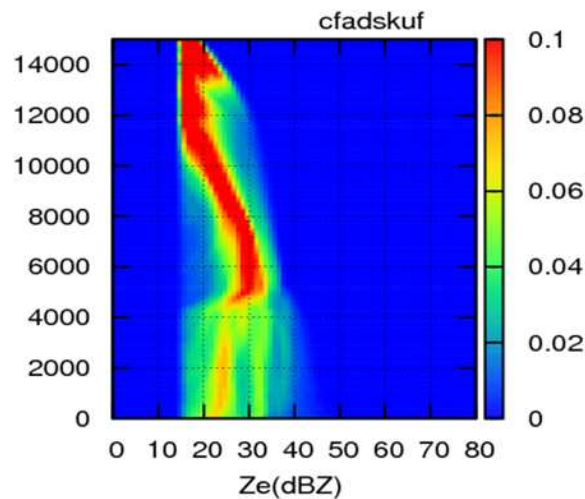
Removed noise using clutter-flag and precipitation-flag



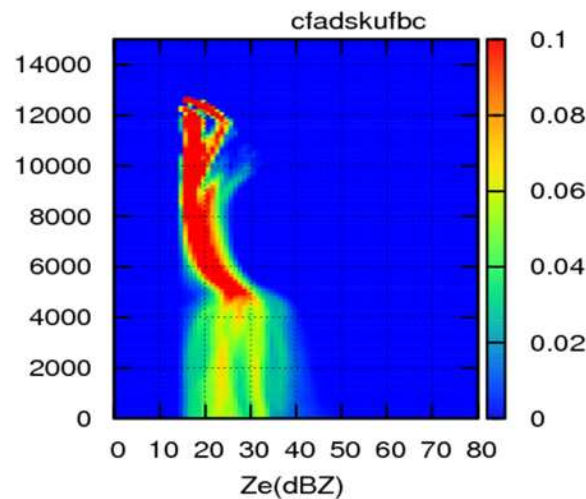
# Quality Control and Bias Correction

- Removal of clutter is very important.
  - flagEcho (by JAXA) can be removed noise almost of all.
- Small noise handling
  - Threshold is defined 15 dBZ.
- Bias of simulated reflectivity handling
  - Bias is removed little by adaptive bias correction.
  - Ice-phase data cannot be assimilated because weak rain ( $< 1$  mm/3h) forecast become negative bias.

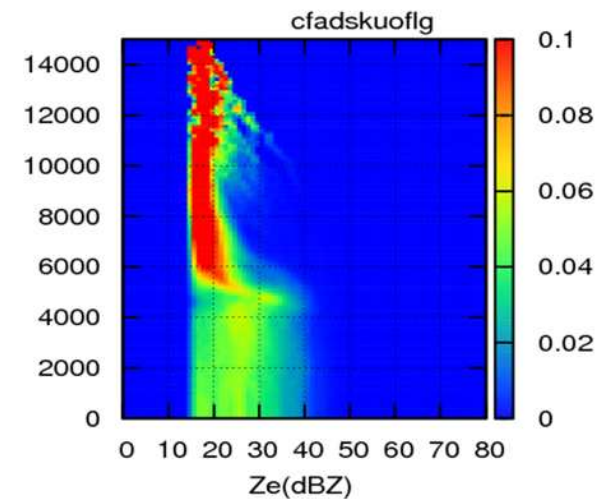
KuPR simulation



KuPR simulation after BC

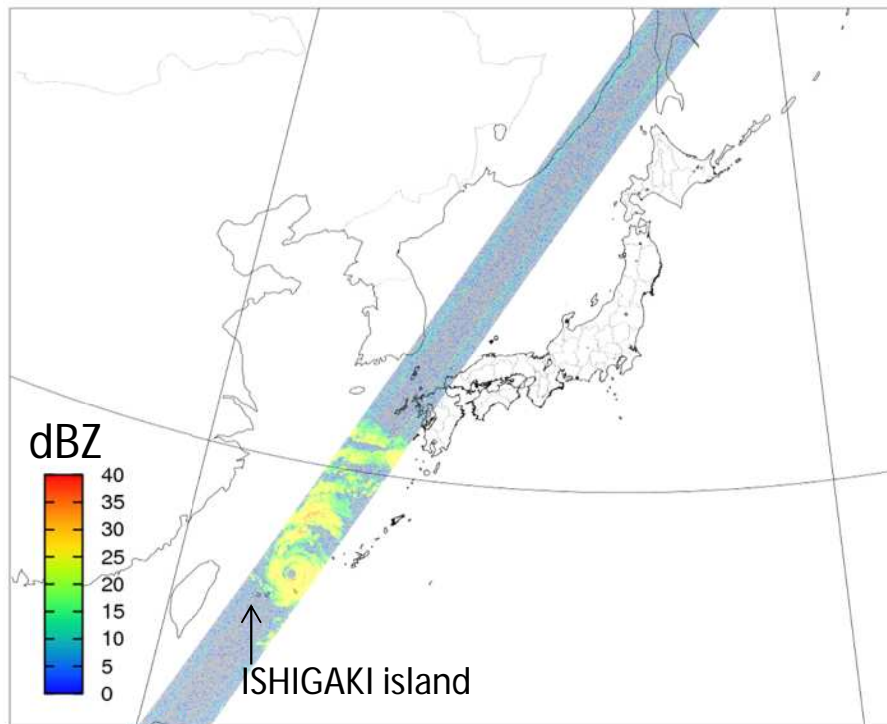


KuPR observaion after QC



# Case study of DPR assimilation

## GPM/DPR KuPR

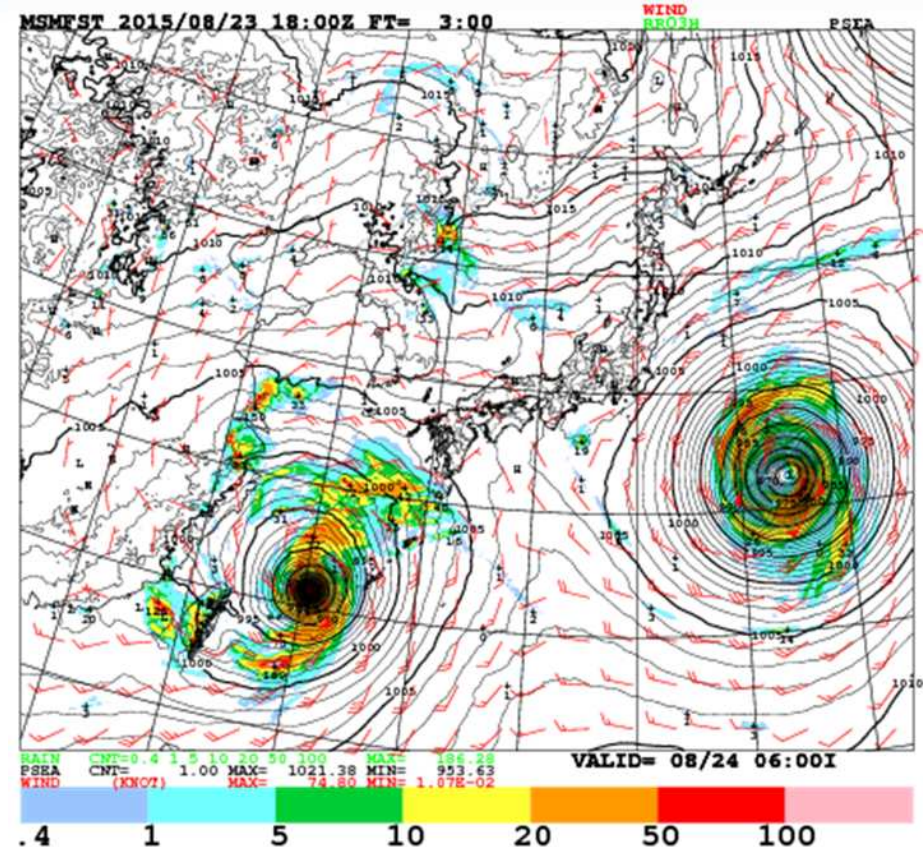


## Typhoon No.15 Goni

Lowest central pressure is 930 hPa.

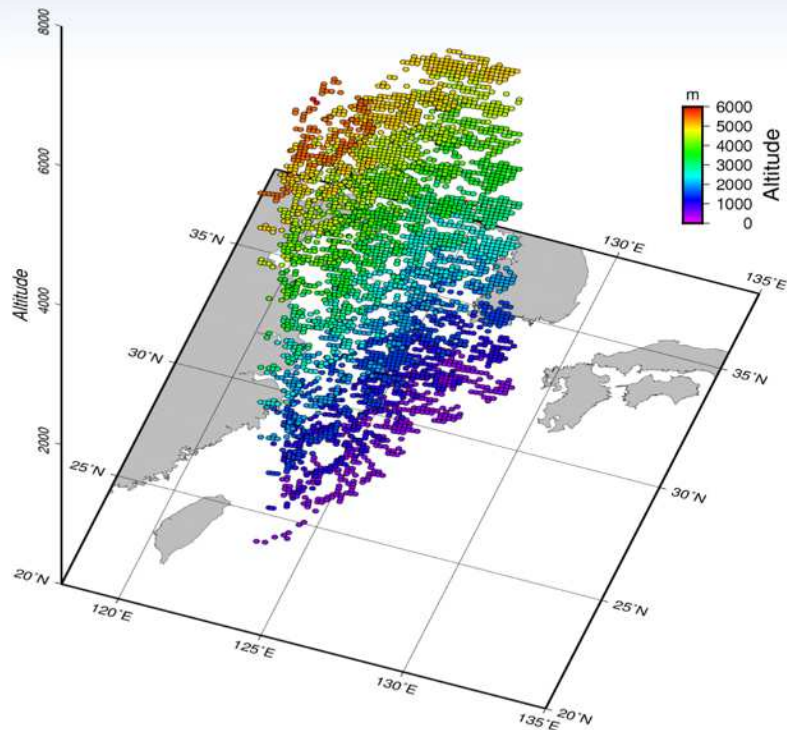
Maximum instantaneous wind speed 71.0 m/s was observed on ISHIGAKI island.

## First Guess of MSM



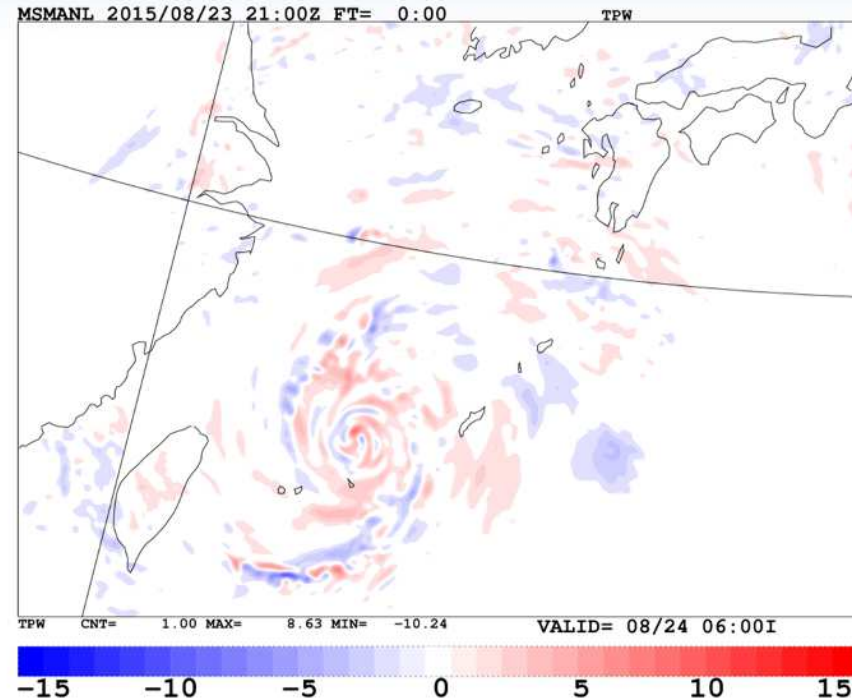
# Case study of DPR assimilation

## Assimilated retrieved RH



Retrieved RH from KuPR and KaPR.  
Only liquid phase data is used, because the ice phase reflectivity has model bias.  
We can assimilate 3-dimensional atmospheric information about moisture !

## Impact of DPR assimilation on initial time



Dry up

Humidify

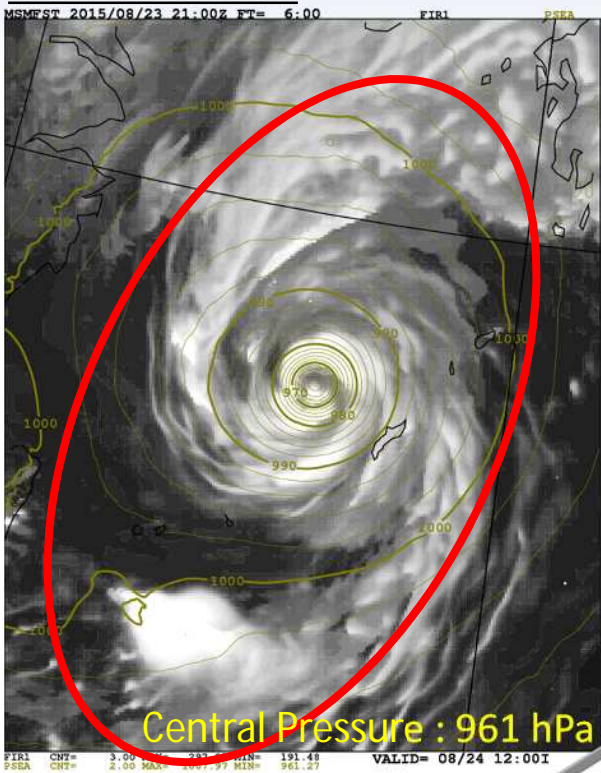
Difference of TPW increment between EXP w/o DPR and EXP w/ DPR.

# Impact of GPM/DPR assimilation

Forecast time 6-hour

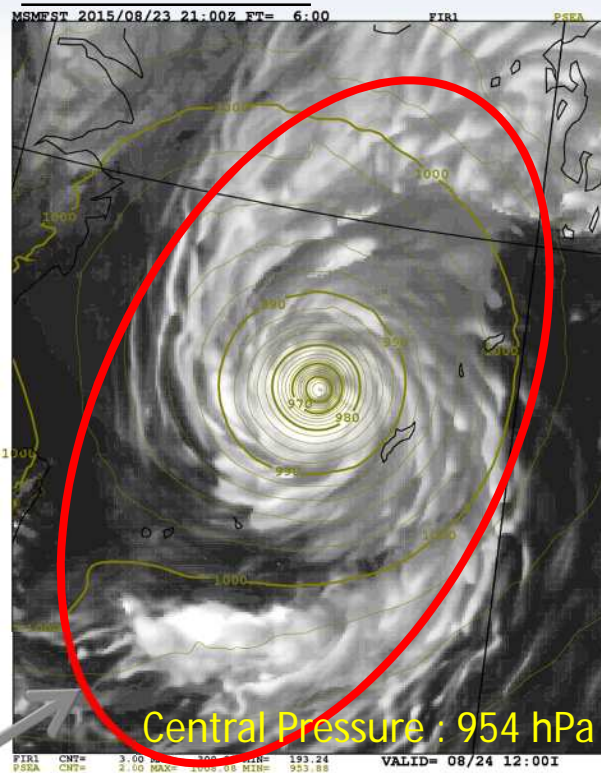
Exp. without GPM/DPR

IR1 Simulation

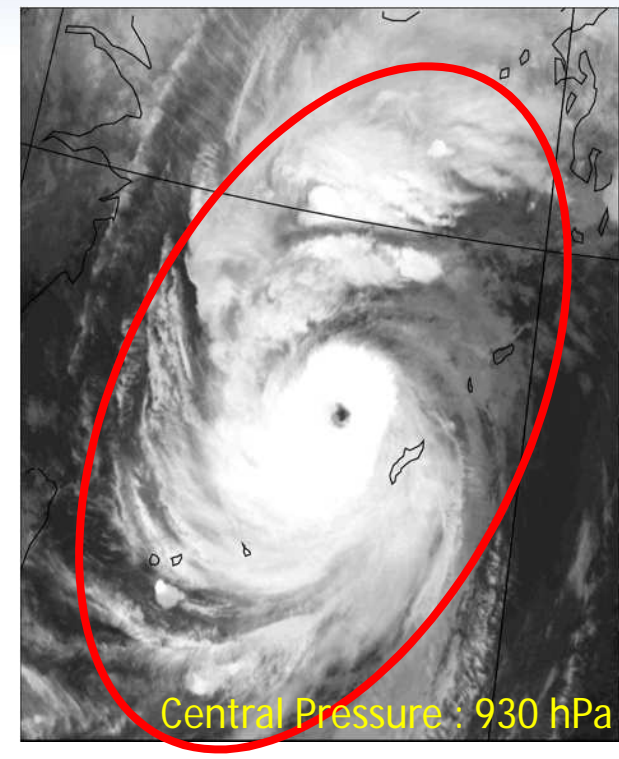


Exp. with GPM/DPR

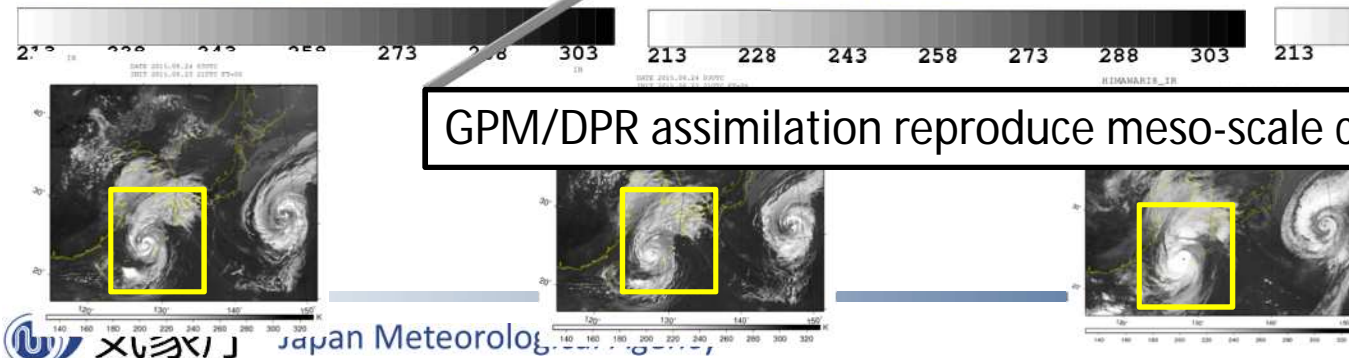
IR1 Simulation



Observation(Himawari-8)



GPM/DPR assimilation reproduce meso-scale convective phenomena.



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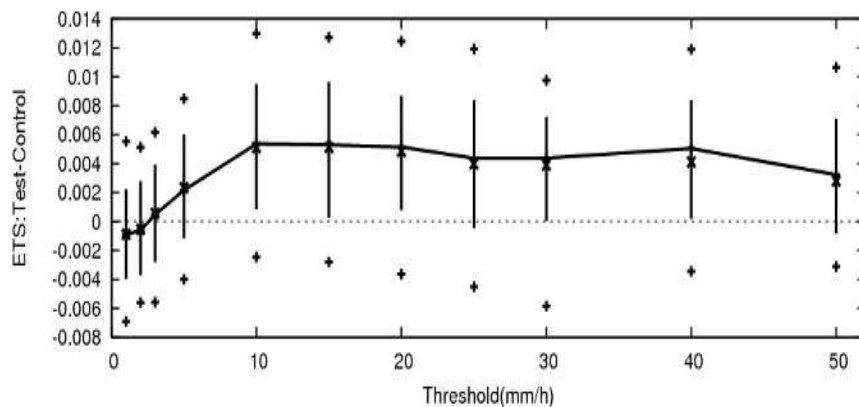
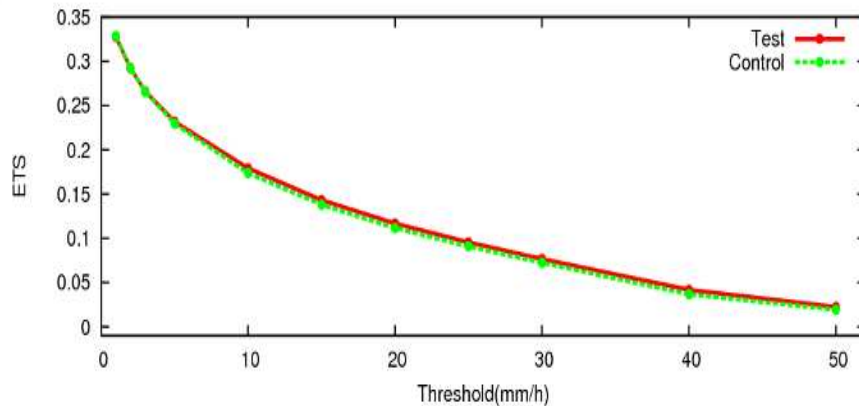
5. Summary

# Performance evaluation test

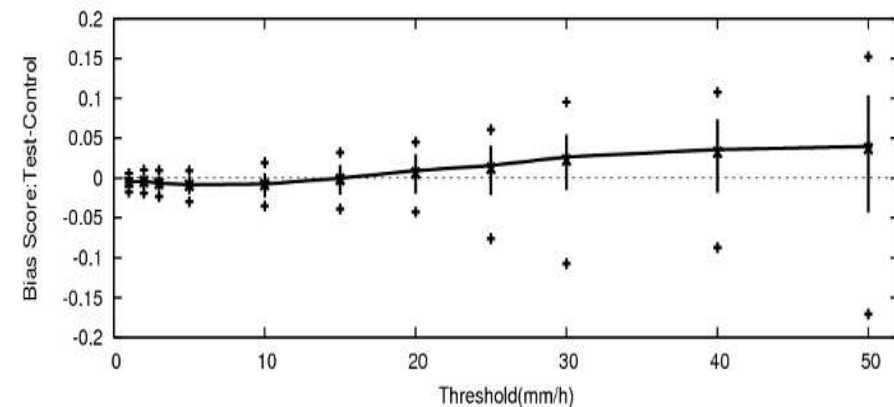
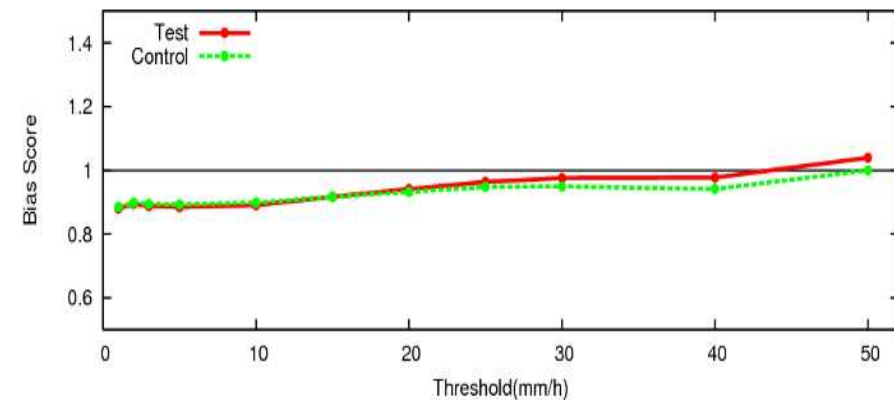
- Meso-scale NWP system
  - Control experiment: with DPR
  - Test experiment: without DPR
    - In both experiments, GMI is not assimilated.
- Experiment periods
  - SUMMER: 7 AUG 2015 ~ 11 Sep 2015
  - WINTER: 10 DEC 2014 ~ 14 DEC 2015

# Equitable Threat Score and Bias Score

## Equitable Threat Score

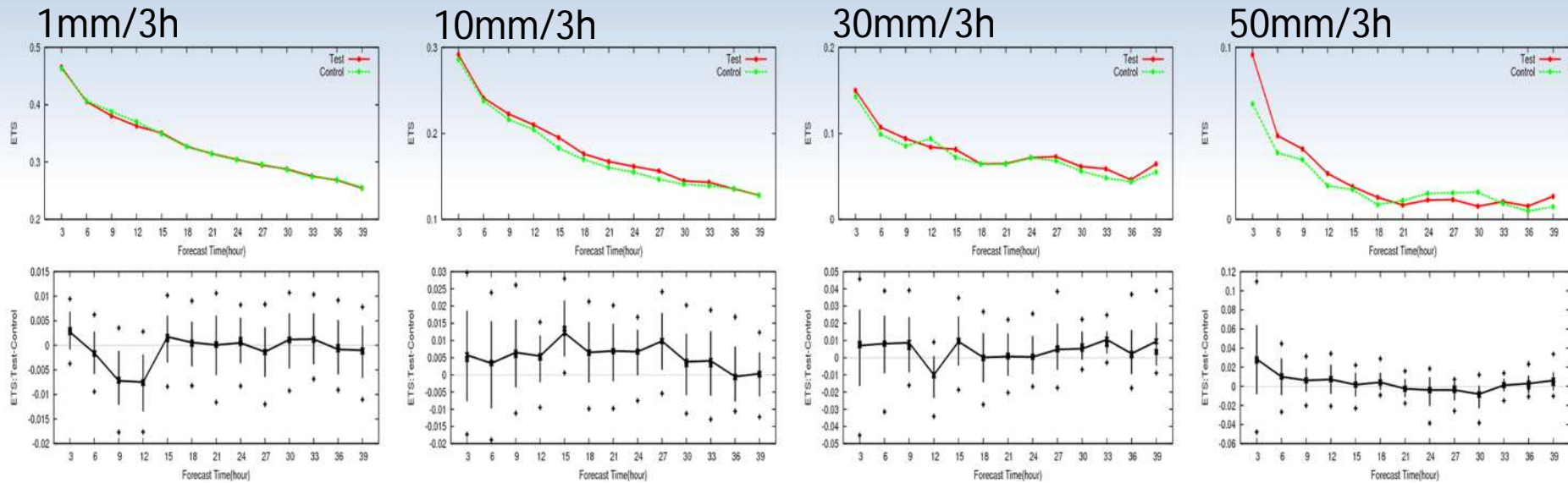


## Bias Score

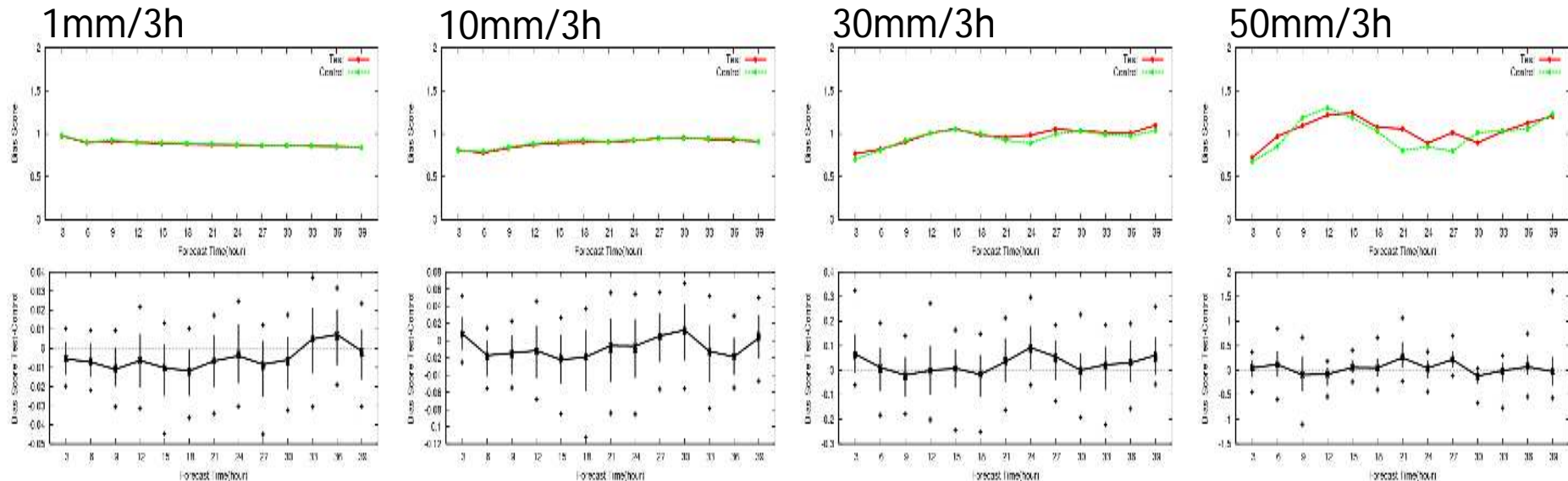


Error bar: 95% confidence interval by block-bootstrap sampling (Wilks 1997)

# Equitable Threat Score



# Bias Score

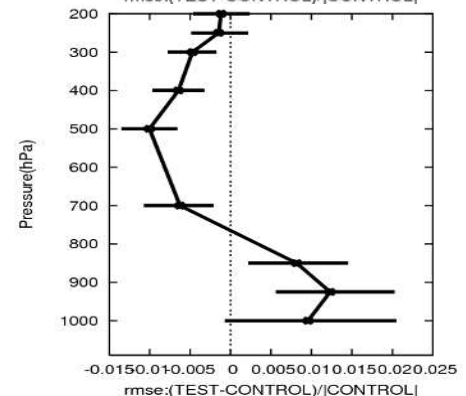
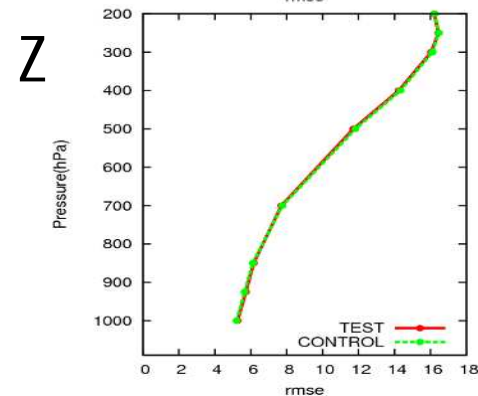
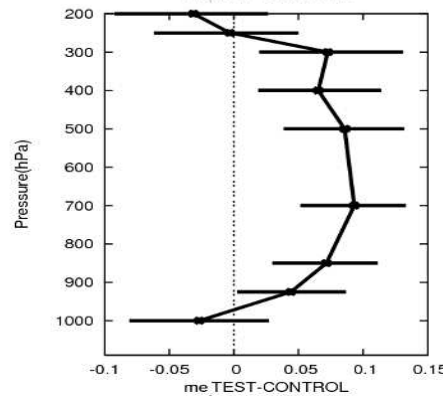
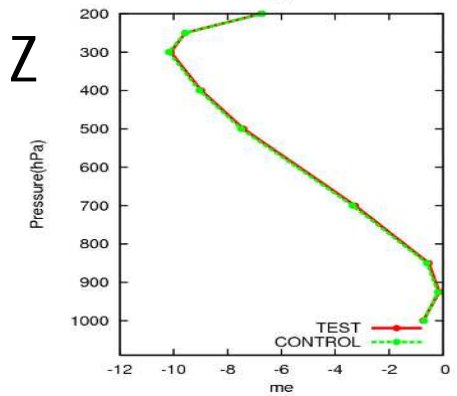
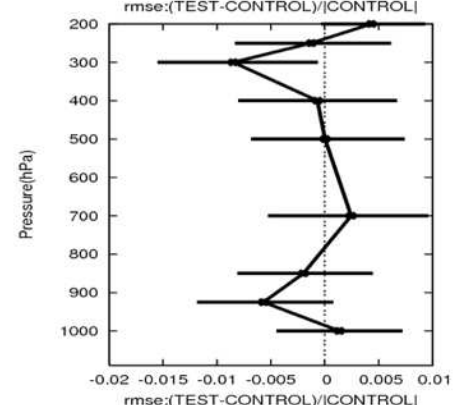
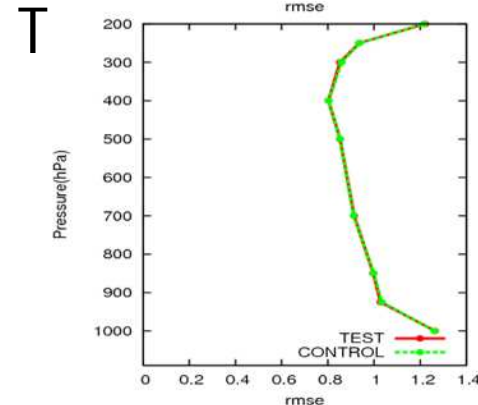
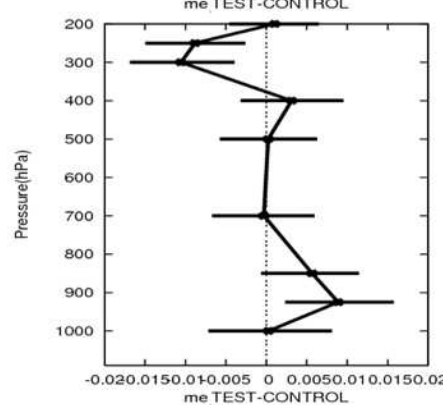
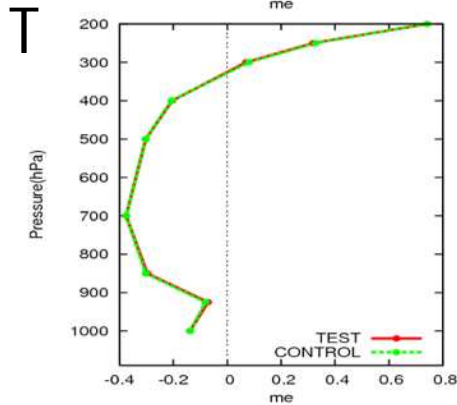
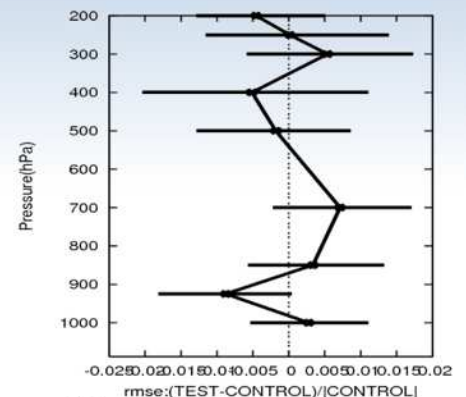
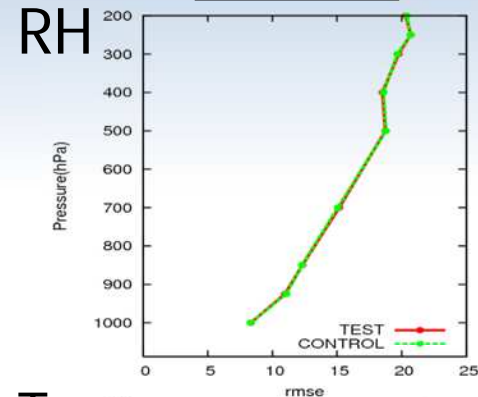
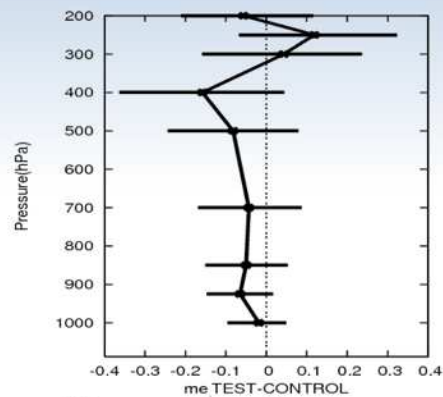
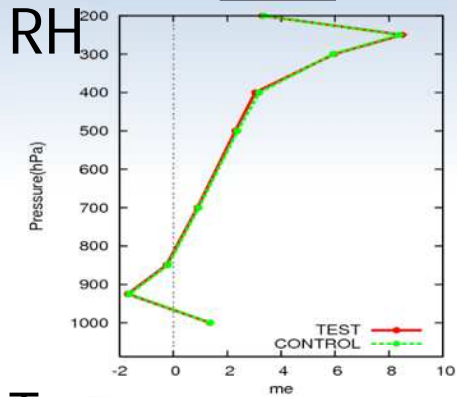




# Sonde verification, Lead time: 0-hour

## ME

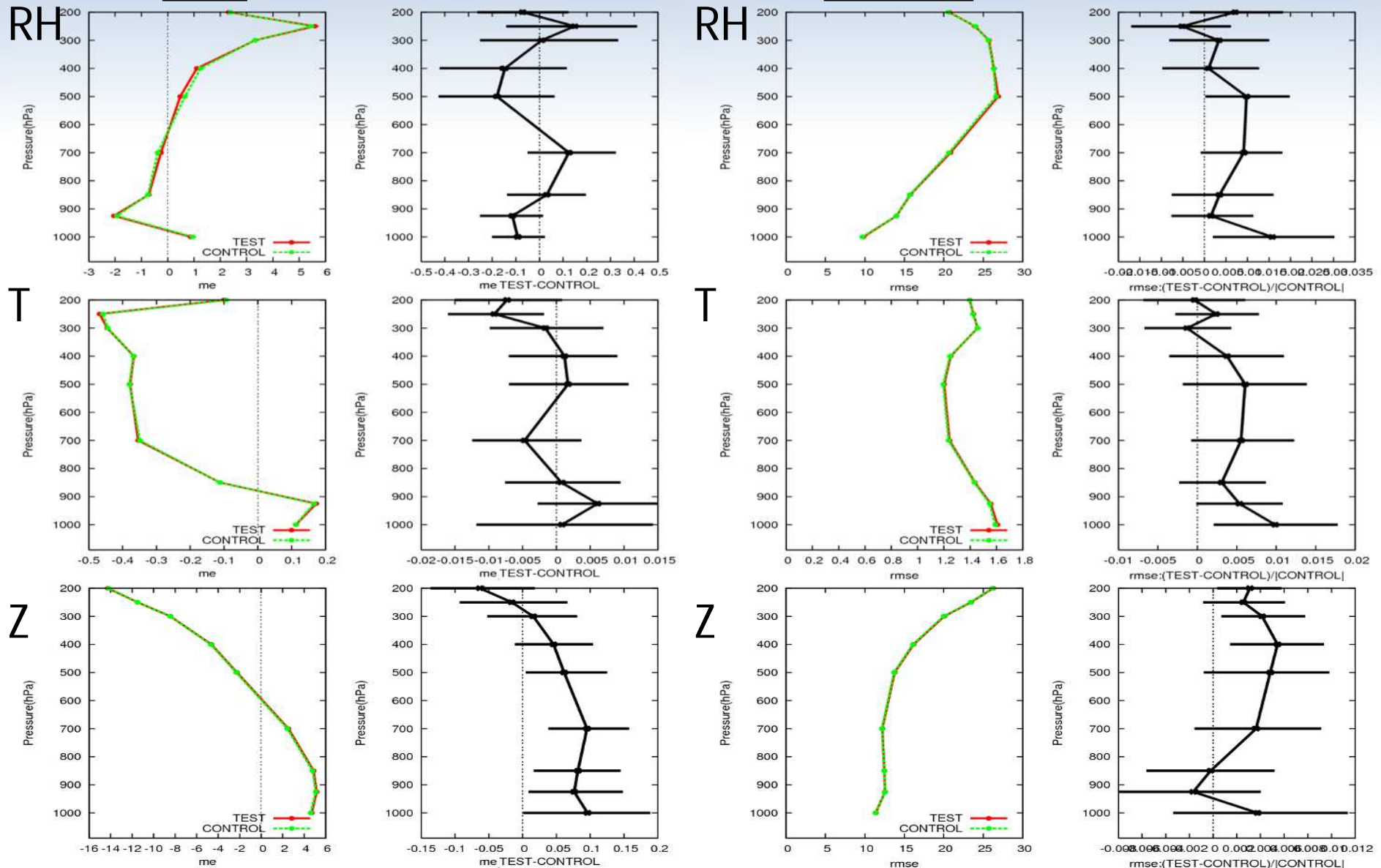
## RMSE



# Sonde verification, Lead time: 39-hour

## ME

## RMSE



# Verification results

- Summer experiment
  - Precipitation
    - ETS over the threshold 10mm/3h is significant improved.
  - Sonde
    - ME of T, Z, RH are improved in initial time.
- Winter experiment
  - Very small impact of DPR
    - Ice phase data are not used.

# Case of Heavy rainfall

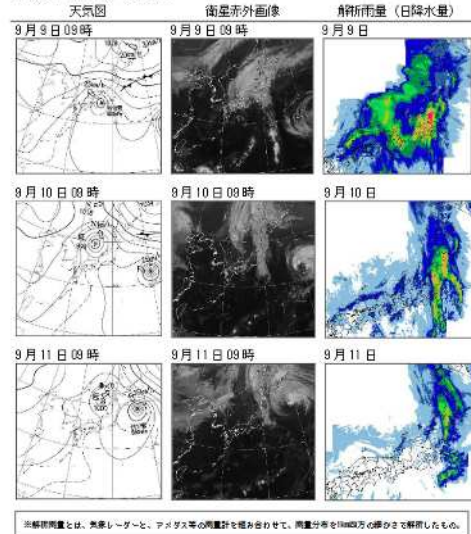
平成27年9月関東・東北豪雨

平成27年9月関東・東北豪雨について

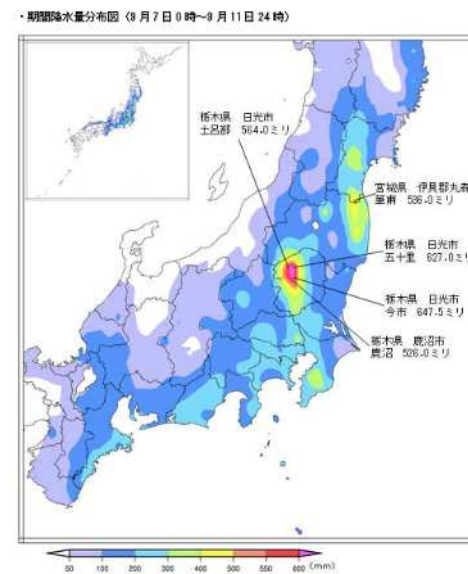
●概要

9月7日03時に  
知多半島に上陸した  
台風第18号が前  
日から11日かけ  
日本の東海上を北上  
が次々と発生した。こ  
9月7日から9月  
丸森町圏まで536.0  
月の月降水量を記録  
けて、栃木県日光市  
では、統計期間が  
を更新するなど、極

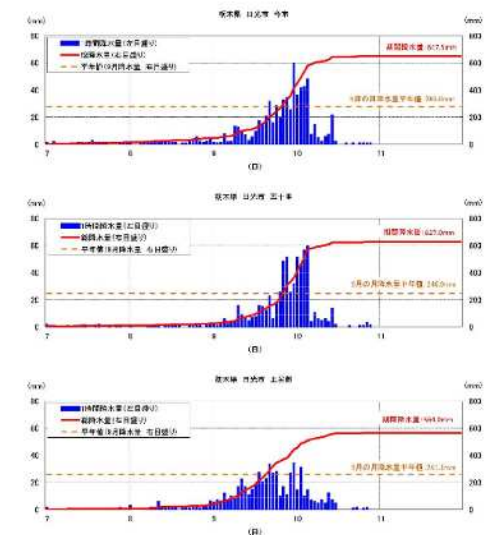
●天気図・衛星画像・解析雨量



●参考：9月7日0時～9月11日24時の観測データ



・降水量時系列図 (9月7日0時～9月11日24時)

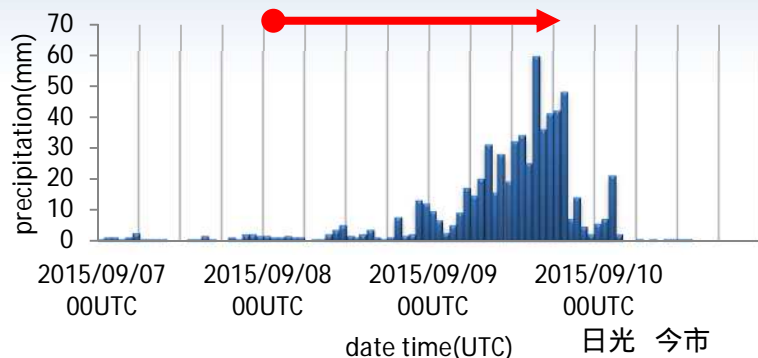


Radar-AMeDAS Precipitation Analysis recorded over 700 mm in 24 hour.

JMA HP: [http://www.jma.go.jp/jma/press/1509/18f/20150918\\_gouumeimei.html](http://www.jma.go.jp/jma/press/1509/18f/20150918_gouumeimei.html)

# Case of Heavy rainfall

Init:2015/09/08 00UTC MSM forecast range

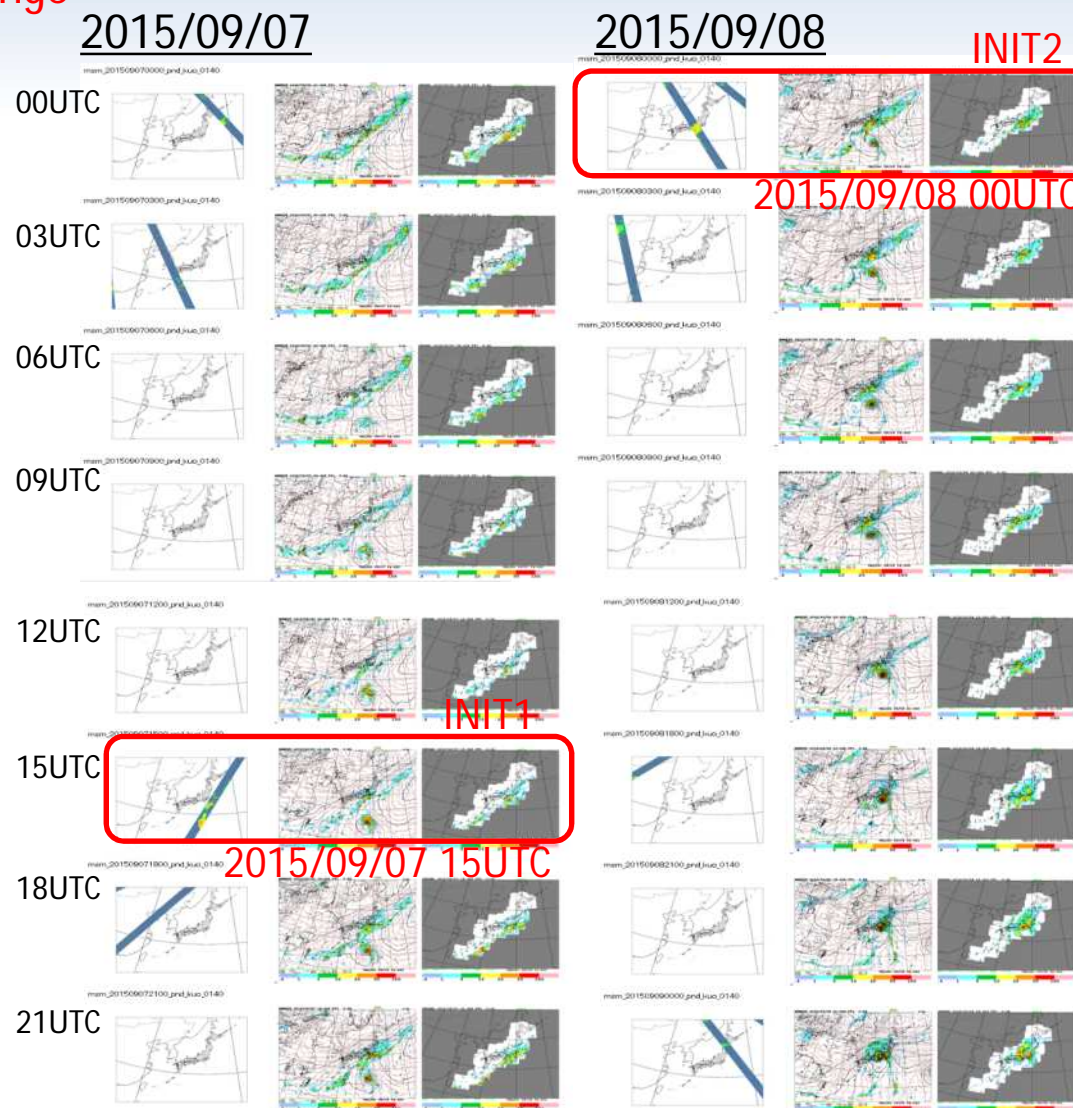


Effect of DPR observation assimilation in to this initial time is carried over by the analysis-forecast cycle.

DPR observation that could contribute to the improvement of heavy rain case, be assimilated in the following initial time.

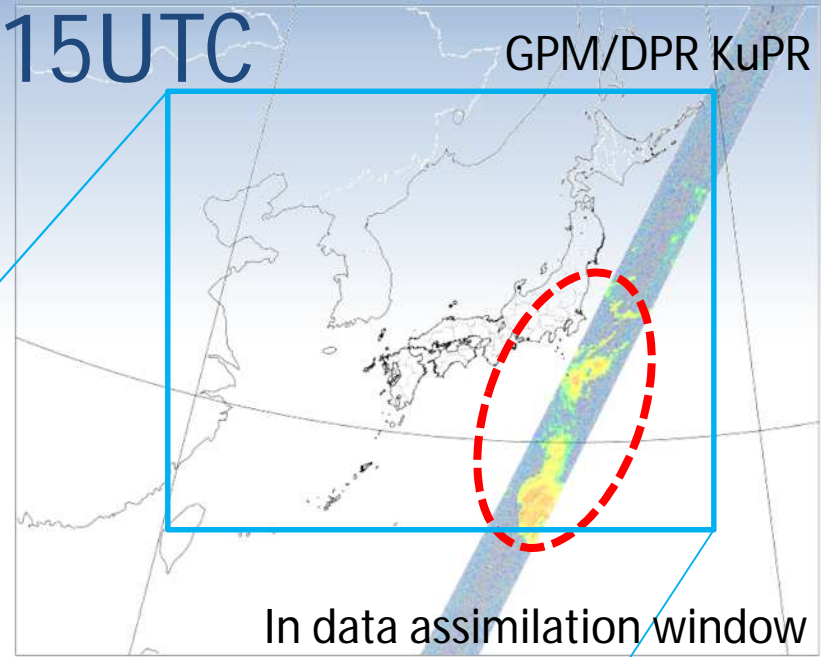
INIT1: 2015/09/07 15UTC

INIT2: 2015/09/08 00UTC

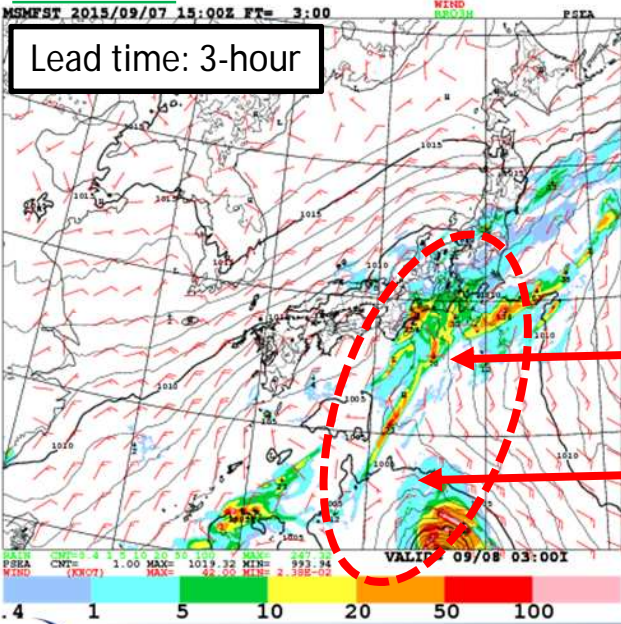


# Case 1 INIT1: 2015-09-07 15UTC

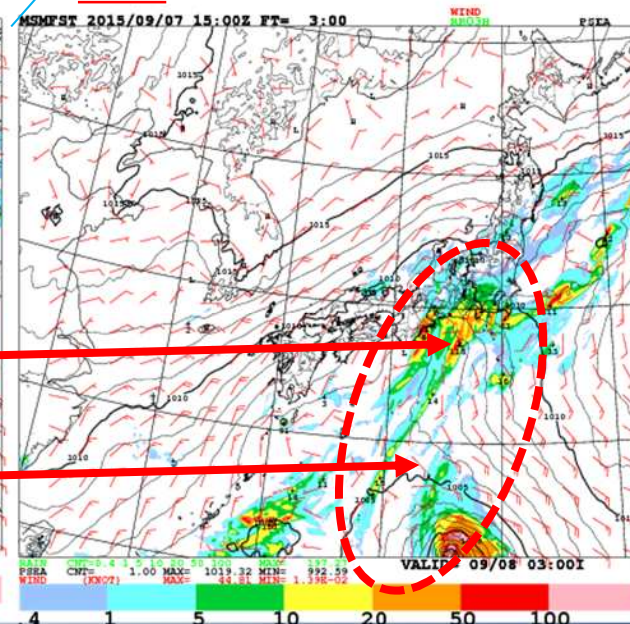
Impact on the rainfall in the south of the sea in the Kanto and Typhoon No. 18. This effect will be taken over to subsequent analysis .



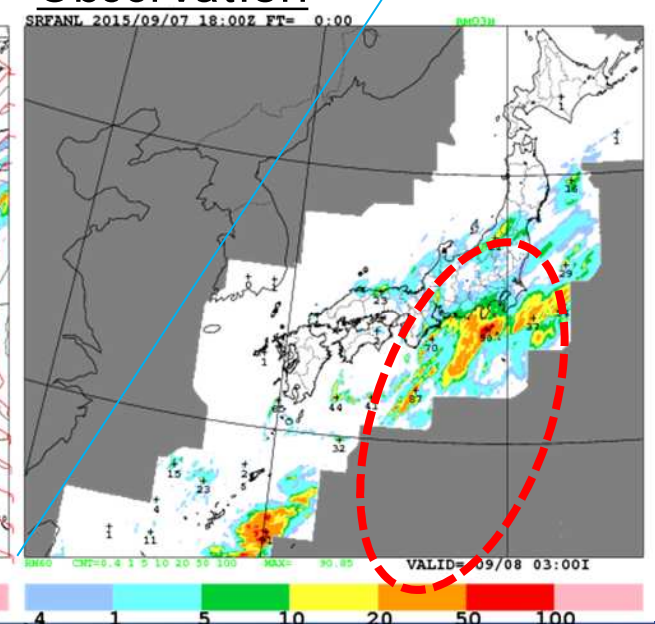
Control



Test



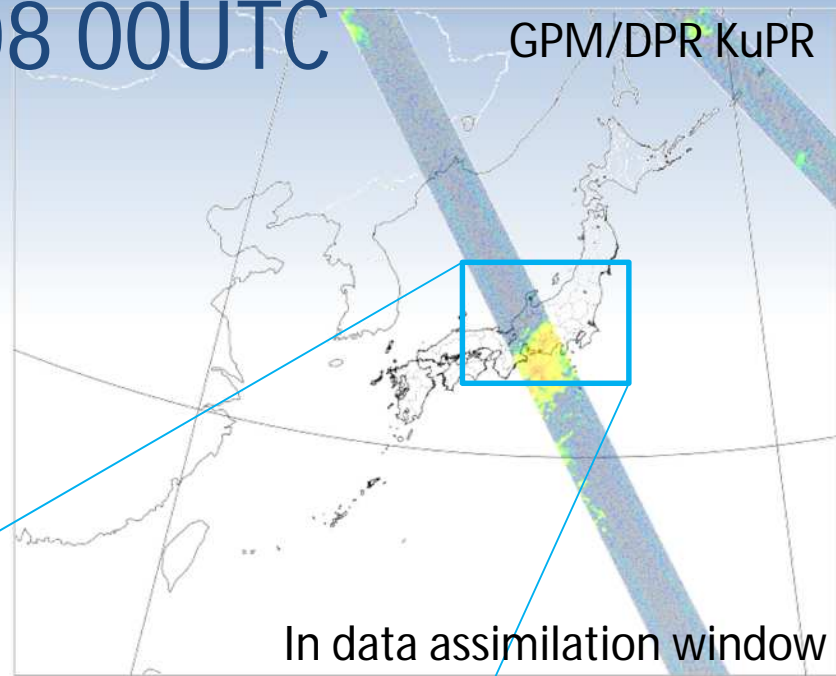
Observation



# Case 2 INIT2: 2015-09-08 00UTC

## lead time: 3-hour

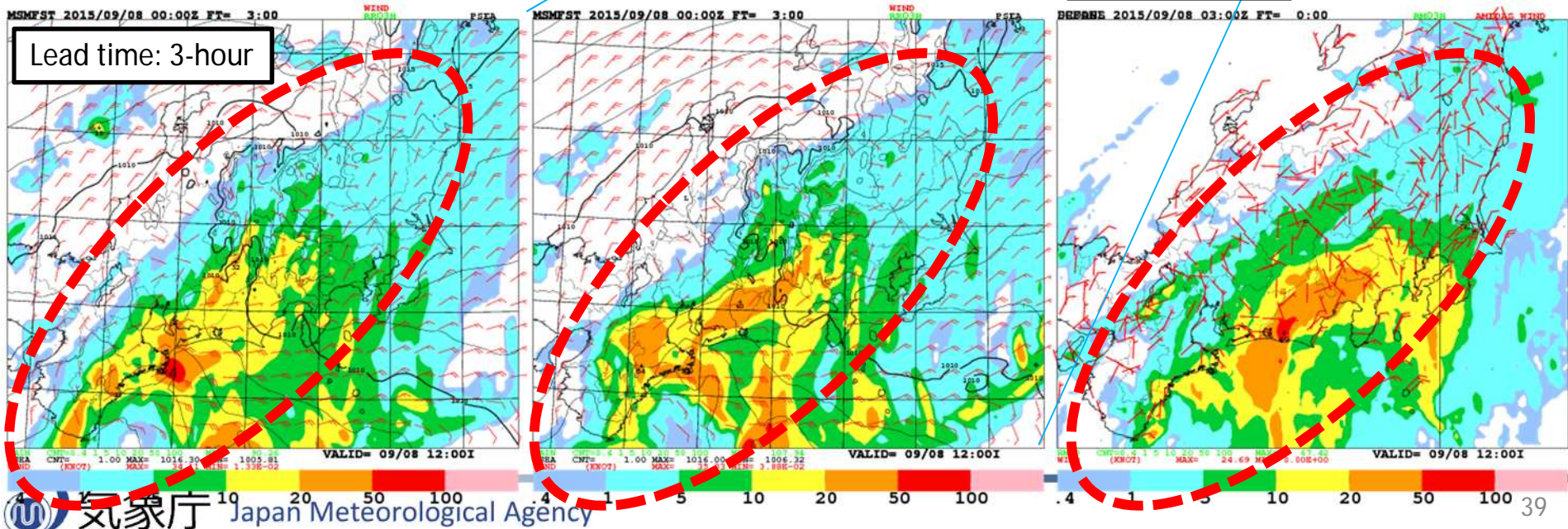
Improvement of precipitation forecast on the DPR path.



Control

Test

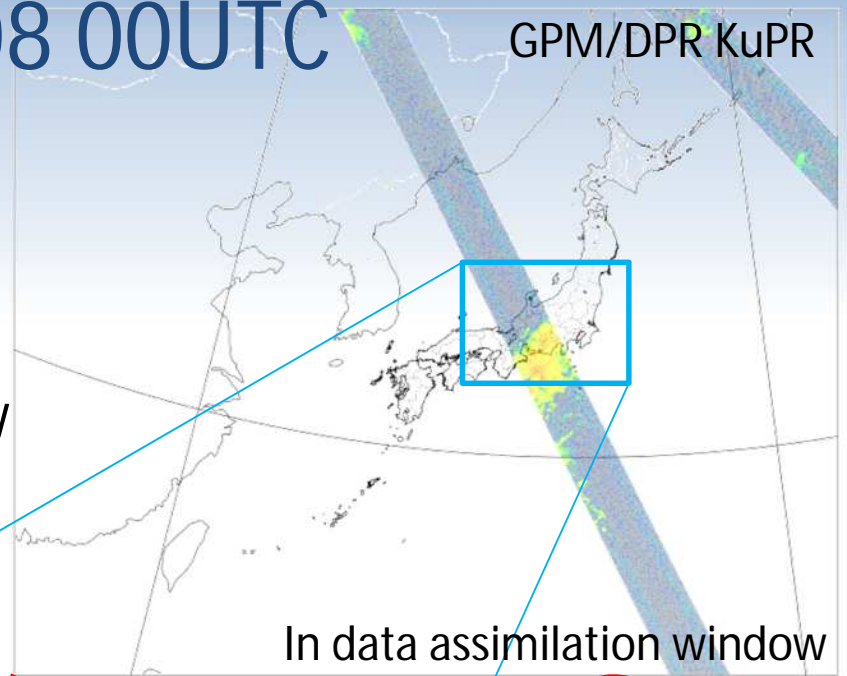
Observation



# Case 2 INIT2: 2015-09-08 00UTC

## lead time: 33-hour

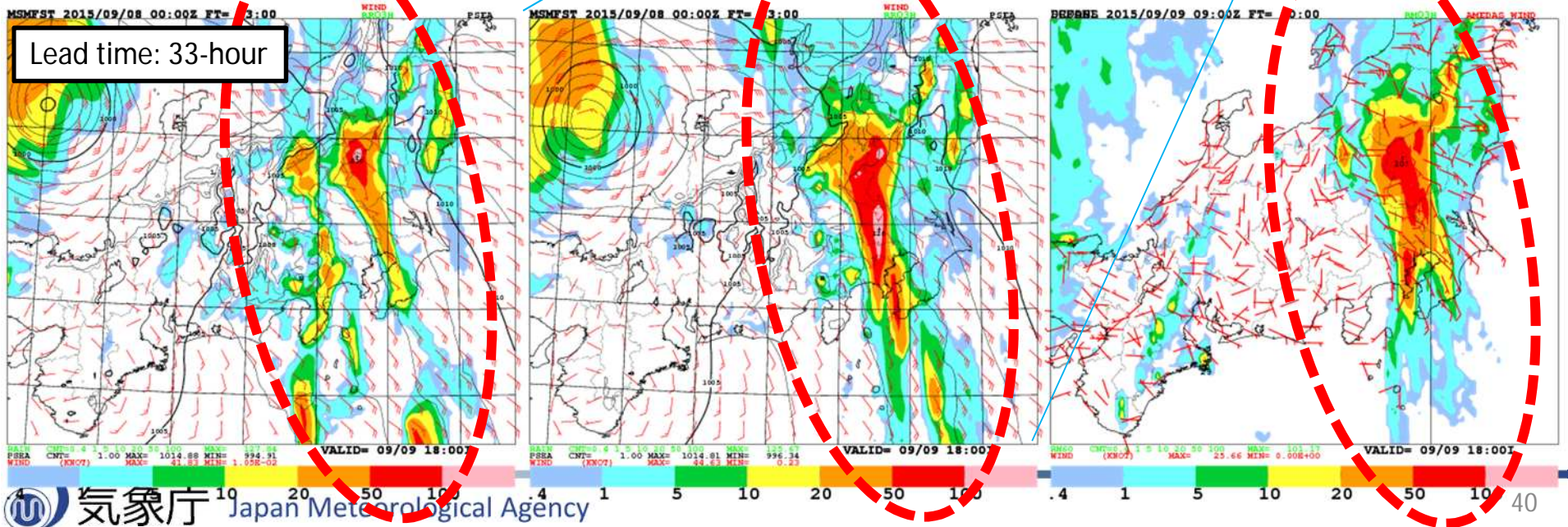
The effect of past assimilation has been taken over in the analysis forecast cycle.  
→ Water vapor in south of Japan has already improved by past analysis(INIT1). Then, in lead time 33h, reproducibility of precipitation is improved also.



Control

Test

Observation





# OUTLINE

1. Introduction

2. Operational NWP system at JMA

3. DPR data assimilation

- Radar simulator

- 1D+4DVAR

  - Relative humidity estimation

- Quality control

4. Performance evaluation

5. Summary

# Summary

- Operational assimilation of GPM/DPR started in March 2016 at JMA.
- Benefit of GPM/DPR data assimilation
  - 3-D information of GPM/DPR is valuable and important data to make initial condition of the meso-scale model.
  - GPM/DPR assimilation improved the forecast of meso-scale convection around Typhoon.
- GPM data will be indispensable data in JMA NWP system.

# Next step of DPR assimilation

## Indirect assimilation

using retrieved RH profiles  
in traditional 4DVAR



## Direct assimilation

using reflectivity (KuPR, KaPR) profiles  
in new Hybrid-4DVAR



THANK YOU FOR YOUR ATTENTION