

# Narrowing uncertainties in climate projections using data science tools

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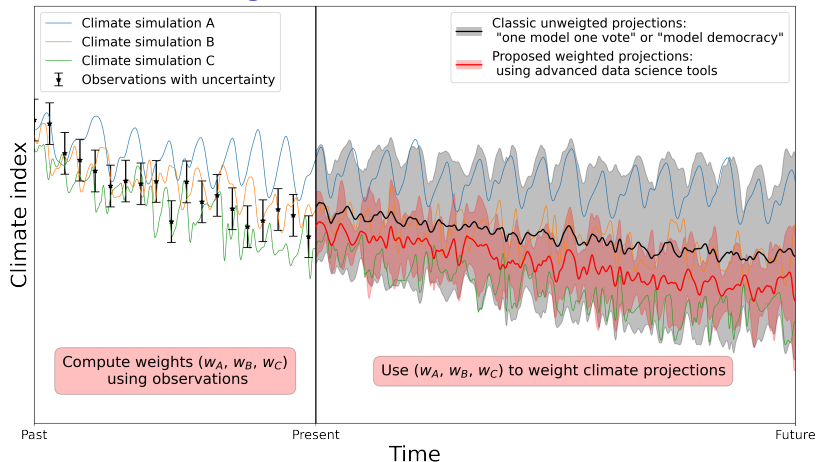
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Univ. Buenos Aires, Argentina<sup>(8)</sup>

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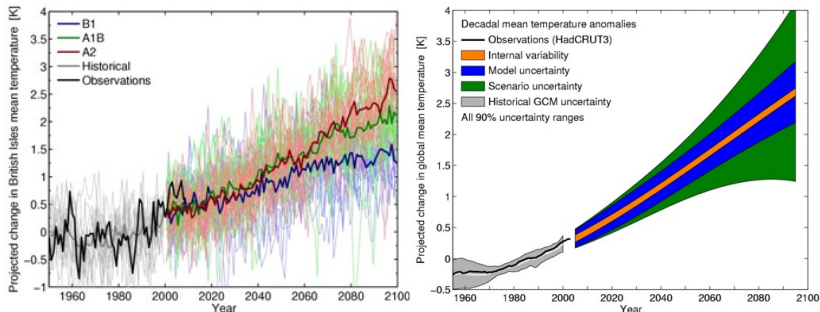
IMT-Atlantique & RIKEN Online Joint Seminar Series

# Context and main goal



- ▶ In the IPCC → ensemble of **unweighted projections** ("one model one vote" or "model democracy", [Knutti, 2010])
- ▶ Idea → **learn weights** from historical observations and simulations, then **propagate weights** to climate projections

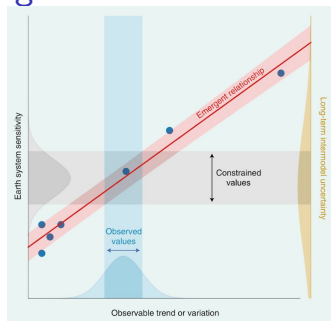
# Climate uncertainties revealed by CMIP



Source: [Hawkins and Sutton, 2011]

- ▶ Climate projections are sensitive to internal, model and scenario uncertainties
- ▶ **Potential to narrow uncertainties**, especially in **regional climate predictions** [Hawkins and Sutton, 2009]

# New at IPCC: emergent constraints



Source: [Eyring et al., 2019]

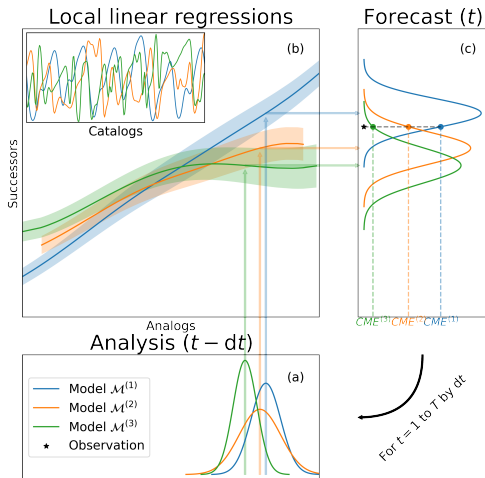
## Pros:

- ▶ Easy to implement (projection using linear regression)
- ▶ Easy to understand (synthetic graphical representation)
- ▶ Do not weight climate simulations (not directly)

## Cons:

- ▶ Causality not obvious (especially for large horizons)
- ▶ Low number of samples to fit the regression
- ▶ Questionable linear relationship and homoscedasticity

# Proposed approach: use advanced data science methods



## Three main steps:

- ▶ (a) **Data assimilation** (ensemble Kalman filter)
- ▶ (b) **Data-driven forecasting** (local linear regression)
- ▶ (c) **Distance obs-forecasts** (contextual model evidence)

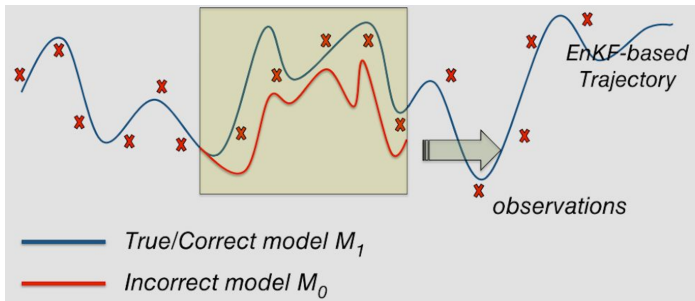
# Ingredient 1: (a) EnKF + (c) contextual model evidence

Contextual model evidence in data assimilation (CME):

$$\mathcal{L}(y(t)|\mathcal{M}_{(i)}) \propto \exp\left(-d_{(i)}(t)^\top \Sigma_{(i)}(t)^{-1}d_{(i)}(t)\right) \quad (1)$$

with the **innovation** defined by its mean and covariance:

$d_{(i)}(t) = y(t) - Hx_{(i)}^f(t)$  and  $\Sigma_{(i)}(t) = HP_{(i)}^f(t)H^\top + R$ .



Source: [Carrassi et al., 2017]

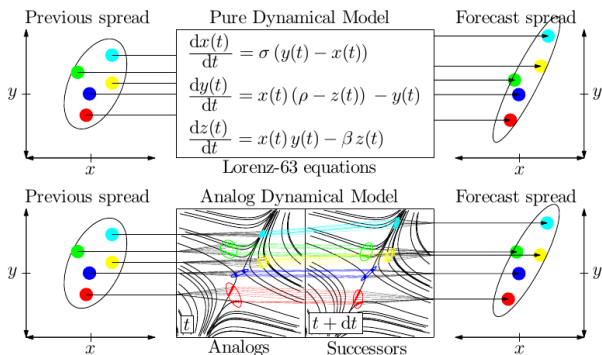
## Ingredient 2: (a) EnKF + (b) analog forecasting

Analog forecasting within data assimilation (AnDA):

$$x(t) = \mathcal{A}(x(t - dt), \eta(t)) \quad (2)$$

$$y(t) = \mathcal{H}(x(t)) + \epsilon(t) \quad (3)$$

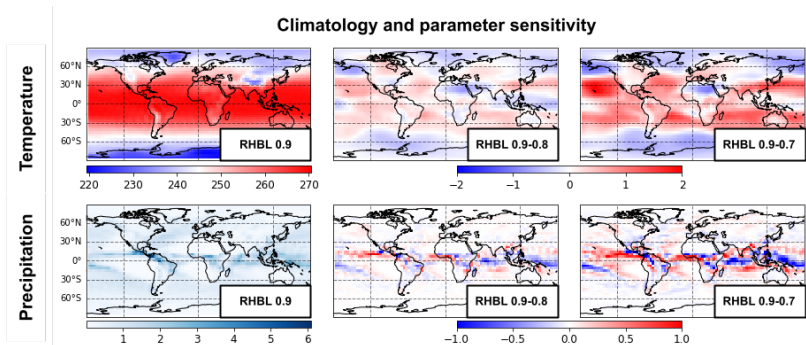
with  $\mathcal{A}$  the **analog forecasting** operator [Lguensat et al., 2017].



Source: [Tandeo et al., 2015]

# Ingredients 1 + 2: (a) EnKF + (b) AF + (c) CME

- ▶ Tested on a **simplified GCM** (SPEEDY, [Molteni, 2003]):
  - ▶ 7 vertical levels,  $96 \times 48$  horizontal grid
  - ▶ simple physics (convection, clouds, radiation, boundary layer)
- ▶ Relative Humidity threshold in the Boundary Layer:
  - ▶ RHBL = 0.9 → the "true" model
  - ▶ RHBL = 0.8 → slightly imperfect model
  - ▶ RHBL = 0.7 → imperfect model

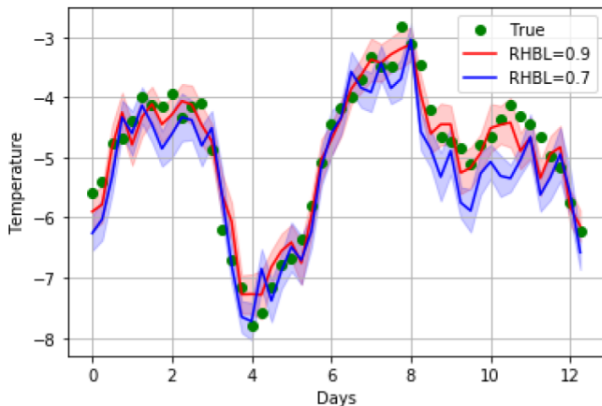




## Ingredients 1 + 2: (a) EnKF + (b) AF + (c) CME

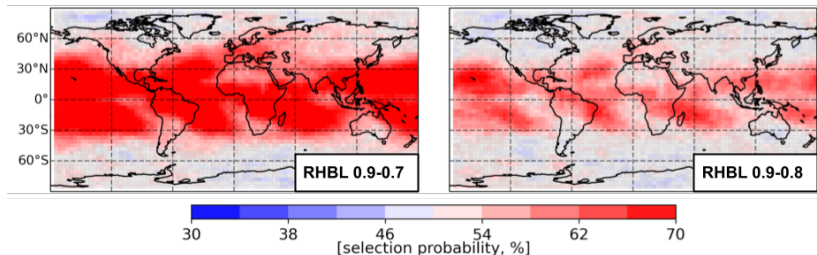
- ▶ Analog data assimilation details:
  - ▶ EnKF with 40 members with adaptive inflation [Miyoshi, 2011]
  - ▶ 30-years catalogs for 3 parameterizations (RHBL 0.9, 0.8, 0.7)
  - ▶ 3D local domains (3 vertical levels,  $3 \times 3$  horizontal grid)
  - ▶ 3 years of noisy observations from RHBL 0.9 (std =  $0.7K$ )

Filtering results using AnDA with 2 catalogs (RHBL=0.9 & RHBL=0.7)



# Ingredients 1 + 2: (a) EnKF + (b) AF + (c) CME

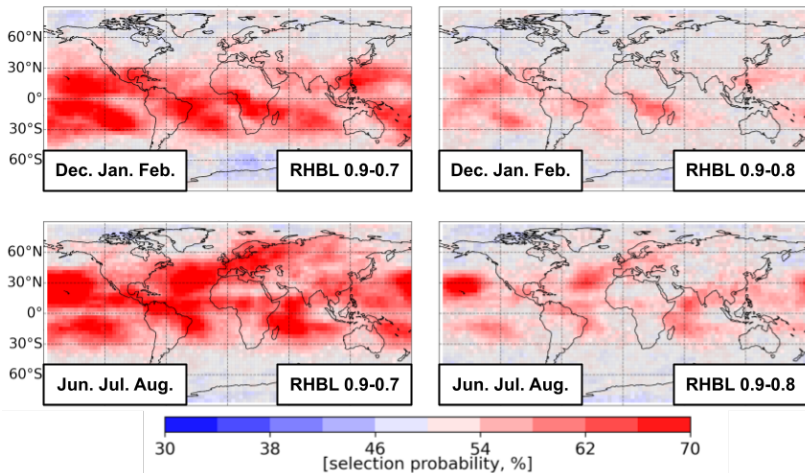
Correct model selection probability for temperature observations



- ▶ Results about model identification (in space):
  - ▶ tropical-subtropical regions affected by model imperfections
  - ▶ degree of imperfection is captured (RHBL 0.7 < 0.8)

# Ingredients 1 + 2: (a) EnKF + (b) AF + (c) CME

Correct model selection probability for temperature observations



- ▶ Results about model identification (in space and time):
  - ▶ sensibility to the RHBL parameter is evolving in time
  - ▶ detection of model imperfection more important in summers (i.e., when there is more convection observed)

# Ingredients 1 + 2: (a) EnKF + (b) AF + (c) CME

## Conclusions:

- ▶ Combination of **advanced data-science methods**
- ▶ Able to compare **short-term model dynamics**
- ▶ Ruiz et al., will be submitted soon to the *Journal of Climate*

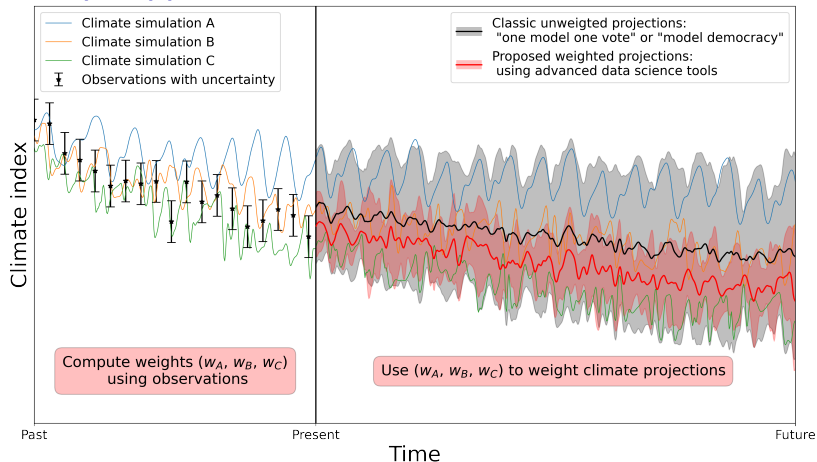
## Pros:

- ▶ **Local approach** (sub-domain, given period, partial variables)
- ▶ **Low-cost procedure** (no need to run climate models)
- ▶ Capture **spatiotemporal differences** in model identifications

## Cons:

- ▶ Need historical numerical simulations
- ▶ Need tuning (analogs, inflation, domain, observations)
- ▶ May seem complicated (but not so much!)

## Next step: application to climate simulations



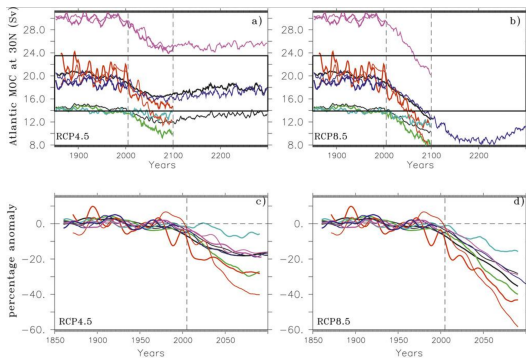
- ▶ Data → compare current observations to CMIP simulations
- ▶ Method → combine data-science methods (DA, AF, CME)
- ▶ Goal 1 → create weighted projections of climate metrics
- ▶ Goal 2 → reduce the uncertainty of climate projections

## Next step: application to climate simulations

Specificity of climate simulations [Knutti et al., 2019]:

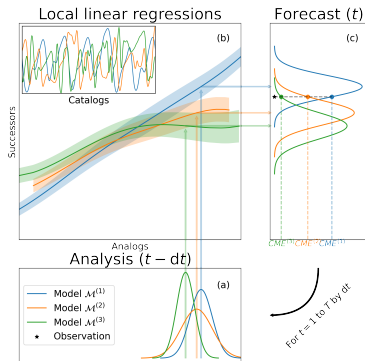
- ▶ **Interdependence** → many CMIP models share ideas, parts of code, or whole components (e.g., the sea ice model)
- ▶ **Performance** → some CMIP models are "good" at representing a specific climate index, other models are not
- ▶ Simulations are sometimes **biased** and need **standardization**

AMOC variations in CMIP5 for 2 RCP scenarios



Source: [Cheng et al., 2013]

## Next step: application to climate simulations



### Caveats and improvement of the methodology:

- ▶ (a) → deal with **model interdependence** (e.g., work with clusters of models), deal with **non-parametric** distributions
- ▶ (b) → find differences in the **short-term dynamics** of climate metrics (especially in the extremes), find **relevant  $dt$**
- ▶ (c) → define more **flexible metrics** (e.g., optimal transport), find **relevant observations** (long time series, knowing noise)

Thank you for your attention! Any questions?

## **MAFALDA:**

**Multi-climate-model Analog Forecasting for  
Attributing Likelihoods using Data Assimilation**



**French ANR program**

**JCJC “young researcher”**

**Under evaluation (2nd round)**





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